

Global changes impacts on water fluxes of the critical zone for the sudanian area in West-Africa

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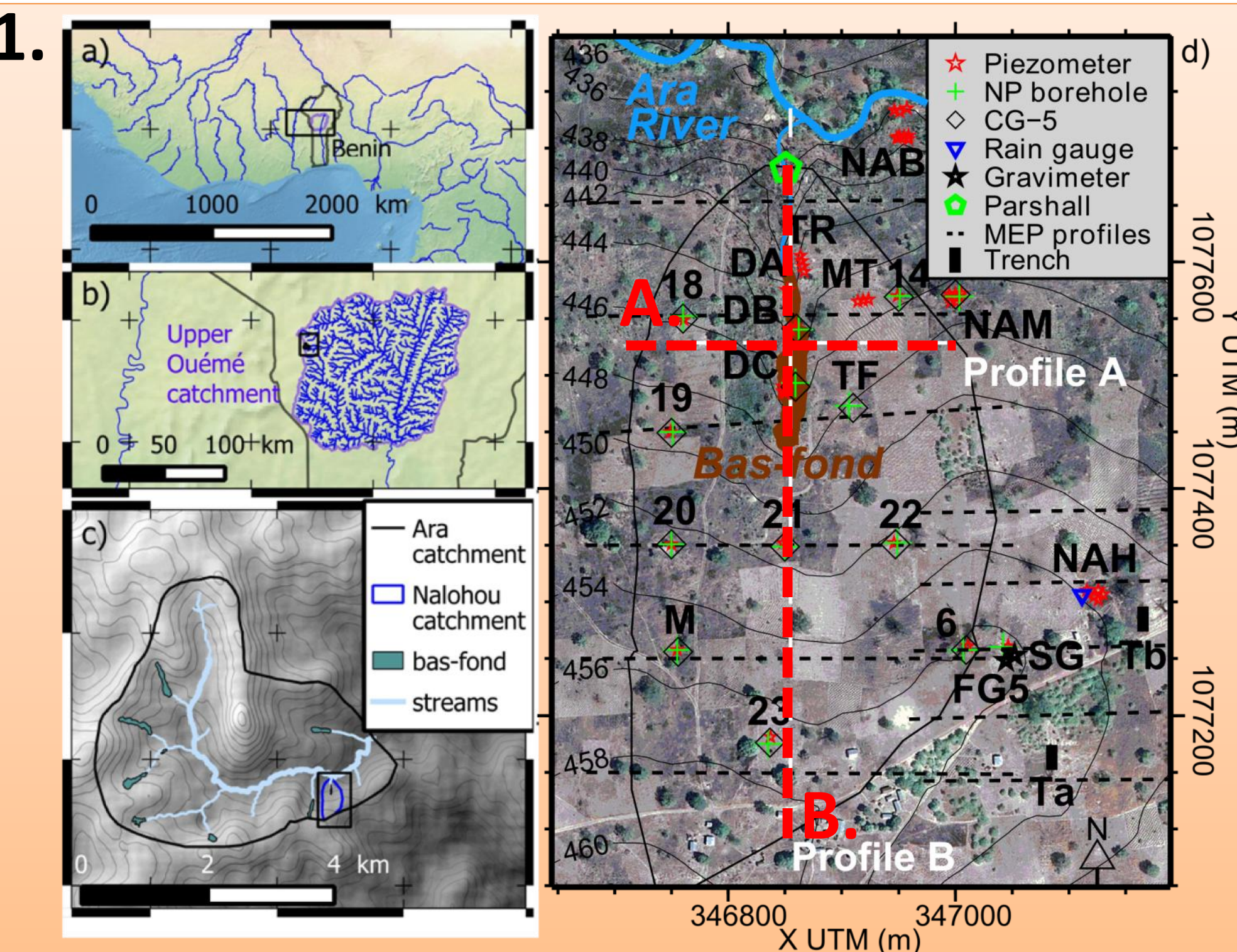
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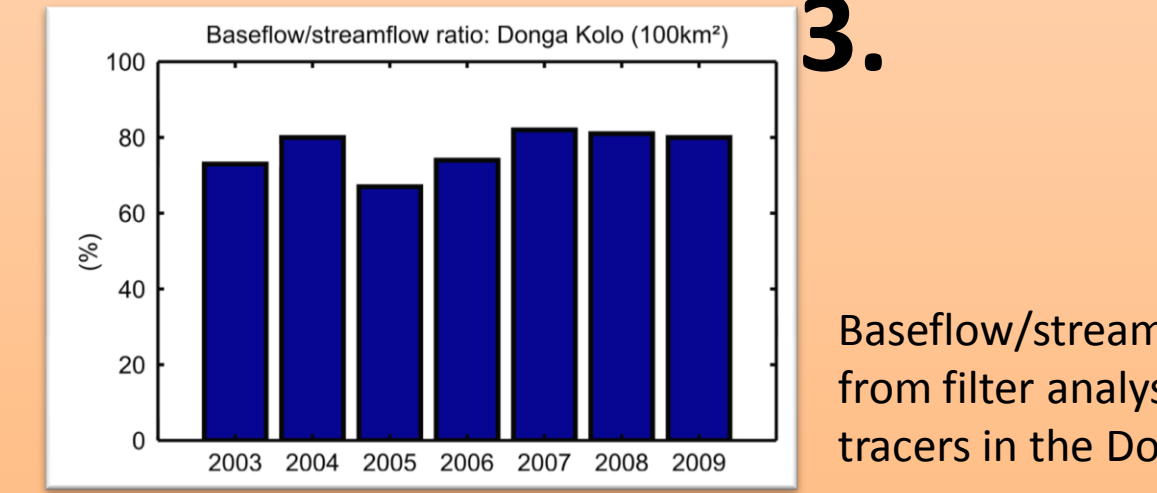
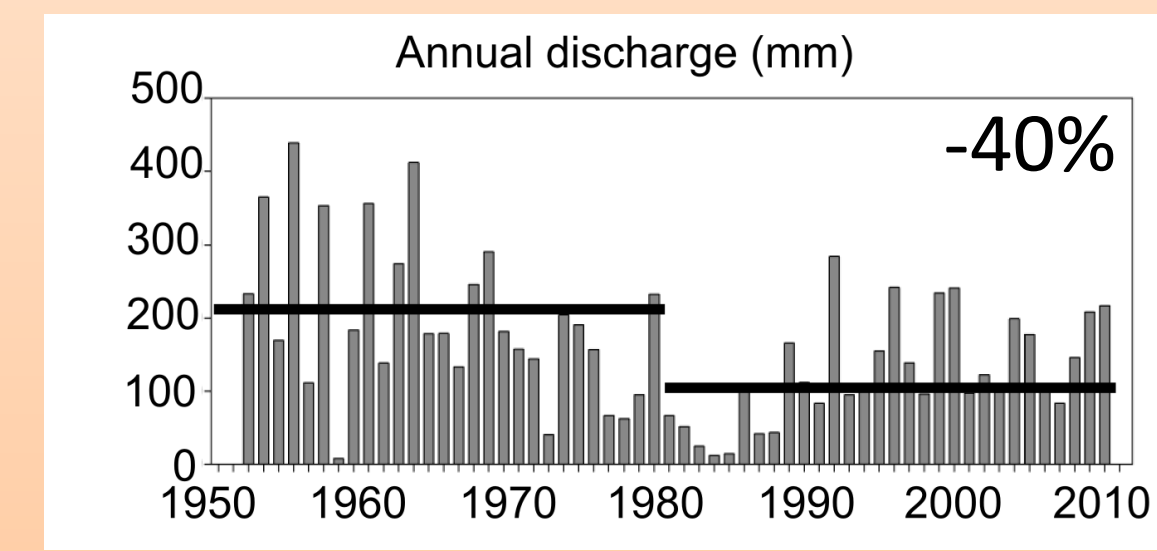
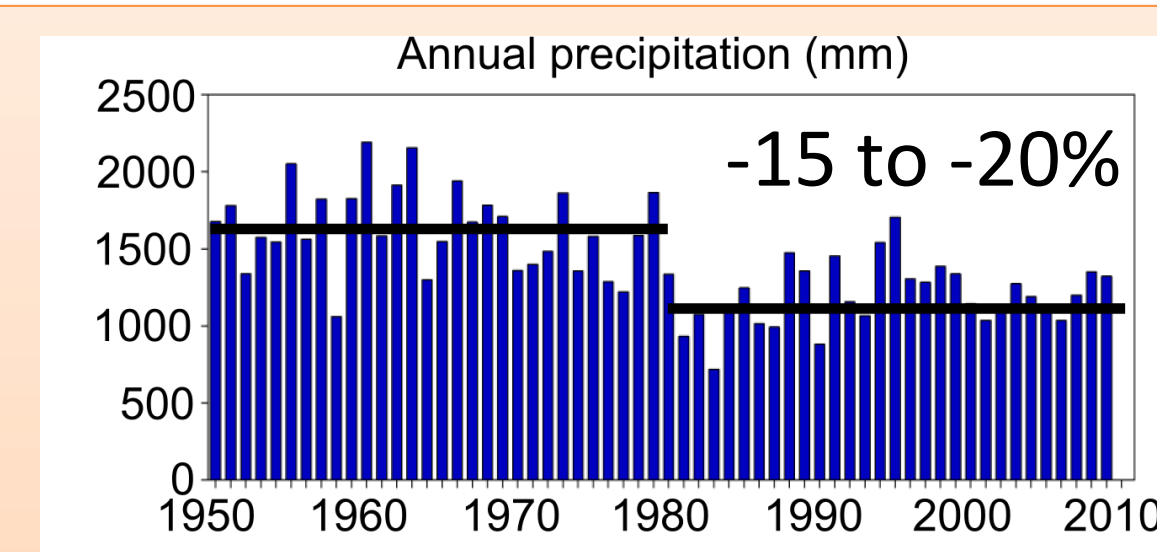
Study area: a 16ha catchment around a seasonally water-logged bas-fond in basement area.

Introduction and context

While in the endoreic Sahel, seasonal streamflow has increased despite the 70'-80' drought and because of simultaneous land-use change, the more humid Sudanian area (1) underwent a **severe streamflow decrease (2)**.

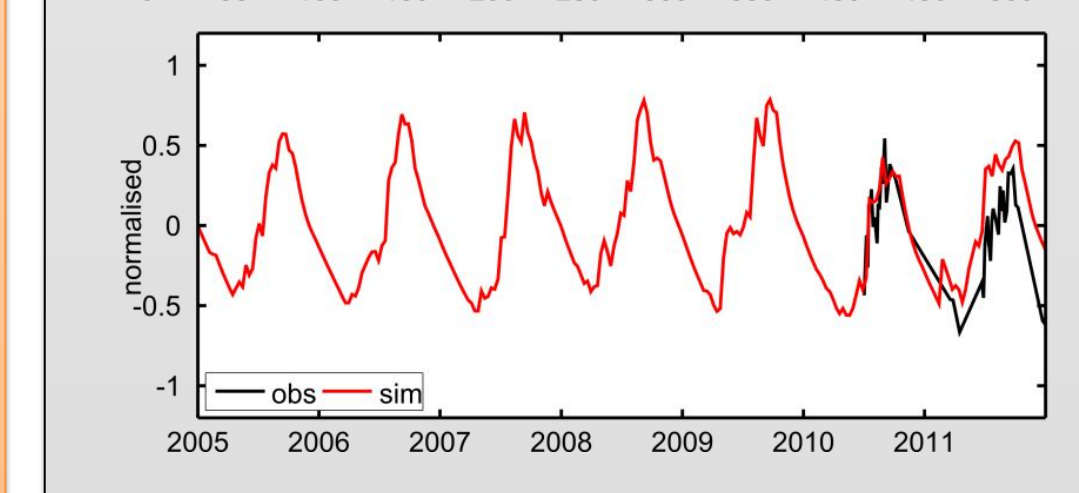
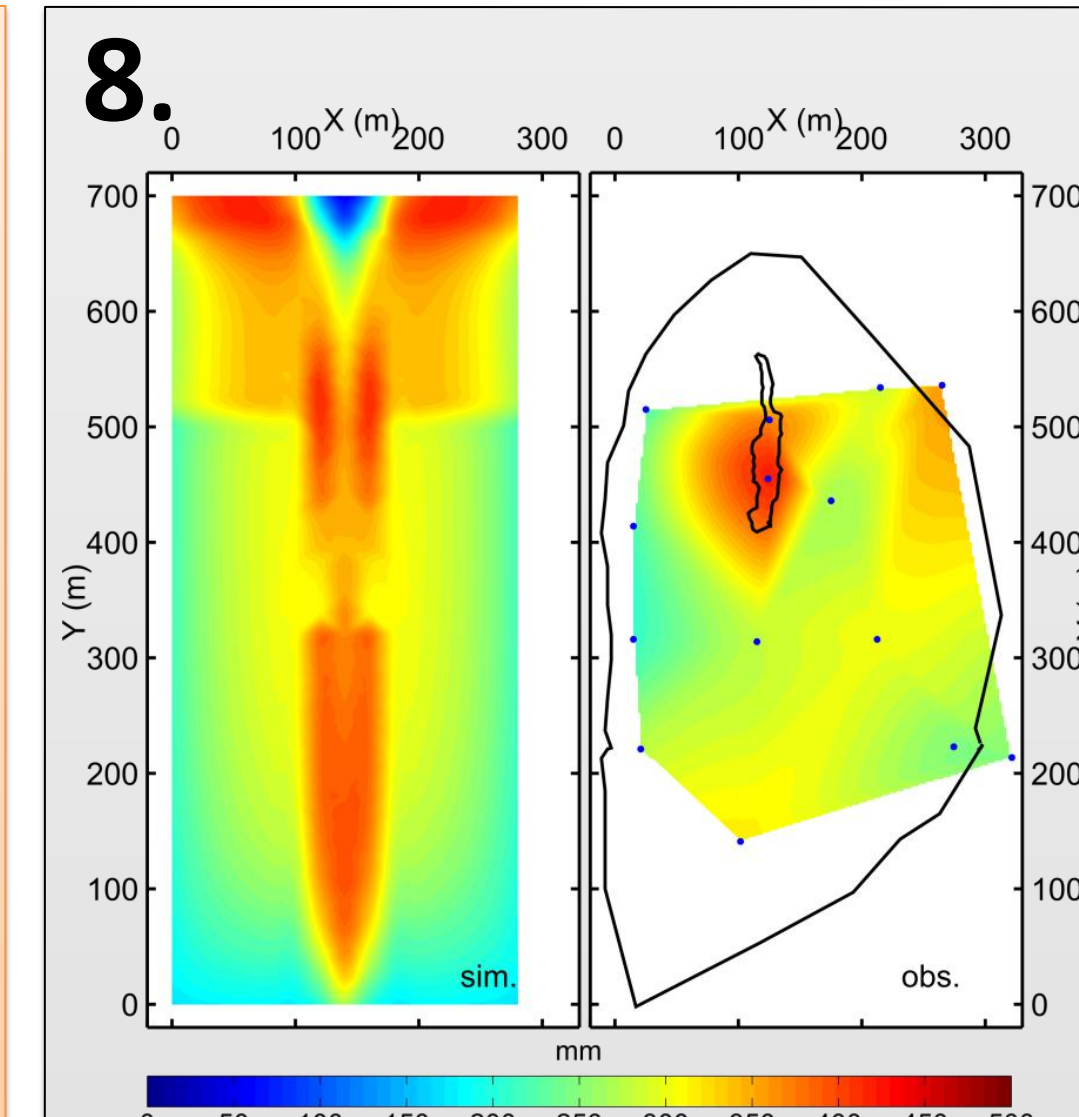
Prevailing processes responsible for such relationships are not fully understood in this area, mostly because **streamflow is dominated by baseflow (3)**, a complex flow component which links all compartment of the critical zone (from bedrock to the top of canopy).

In anticipation of **possible extreme precipitation increase** in the future, as well as **expected land-use and land-cover changes**, precipitation-discharge relationships need to be understood **within a critical zone perspective**.



Precipitation and discharge in the Upper Ouémé valley (14300km²)

Baseflow/streamflow decomposition from filter analysis and geochemical tracers in the Donga Kolo catchment



Simulated (left) and observed (right) seasonal storage amplitude

A. Material & Method

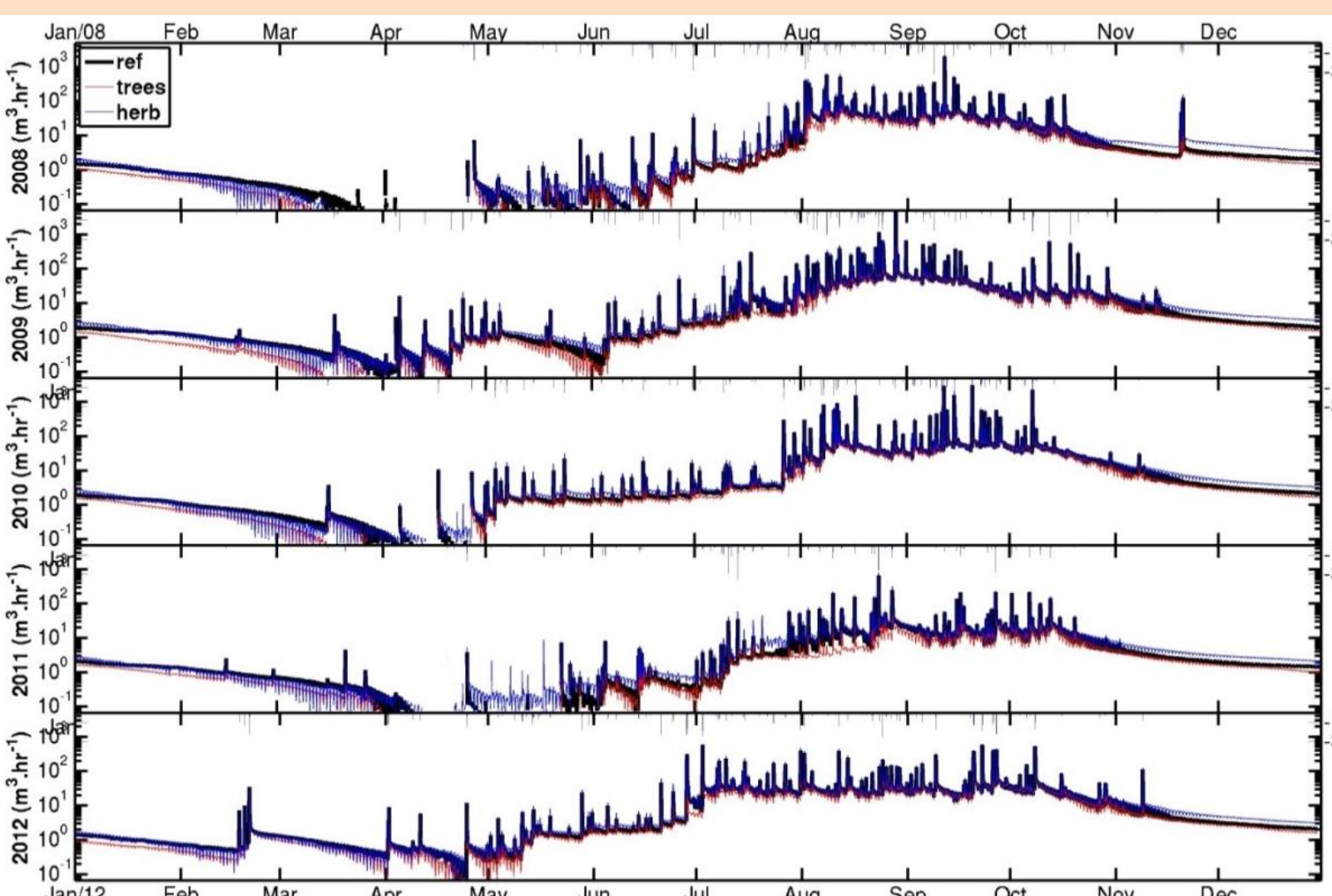
In this study, we aim to:

- present our **conceptual model** of a headwater catchment (16ha), built based on many local observations (1)
- present the **physically-based critical zone model** we use to simulate the hydrological functioning
- Validate a reference simulation of a **simplified V-shaped catchment** with in-situ observations
- test the **sensitivity of the model to land cover**.
- apply the model to the **upper Oueme catchment (14300km²)**, the meso-scale site of the AMMA-CATCH observatory

E. Virtual experiment

- We test two extreme cases: 1) a fully tree-covered catchment and 2) an herbaceous-covered catchment.
- **ET differences** between the two cases vary **between 6 and 13% of yearly Precipitation**.
- **30% less streamflow with tree covers than with herbaceous covers (11)**.

Year	P(mm) [%]	ET (mm) [%]		Q (mm) [%]		S (mm) [%]	
		Ref.	trees	Ref.	trees	Ref.	trees
2006	913 [100]	762 [83]	897 [98]	700 [77]	174 [19]	121 [13]	206 [23]
2007	1214 [100]	828 [68]	914 [75]	788 [65]	362 [30]	281 [23]	405 [33]
2008	1211 [100]	789 [65]	900 [74]	738 [61]	392 [32]	305 [25]	442 [37]
2009	1496 [100]	957 [64]	1038 [69]	919 [61]	565 [38]	459 [31]	613 [41]
2010	1572 [100]	892 [58]	963 [63]	861 [56]	644 [42]	564 [37]	683 [45]
2011	934 [100]	769 [82]	893 [96]	714 [76]	216 [23]	141 [15]	255 [27]
2012	1422 [100]	862 [61]	929 [65]	834 [59]	512 [36]	419 [29]	551 [39]

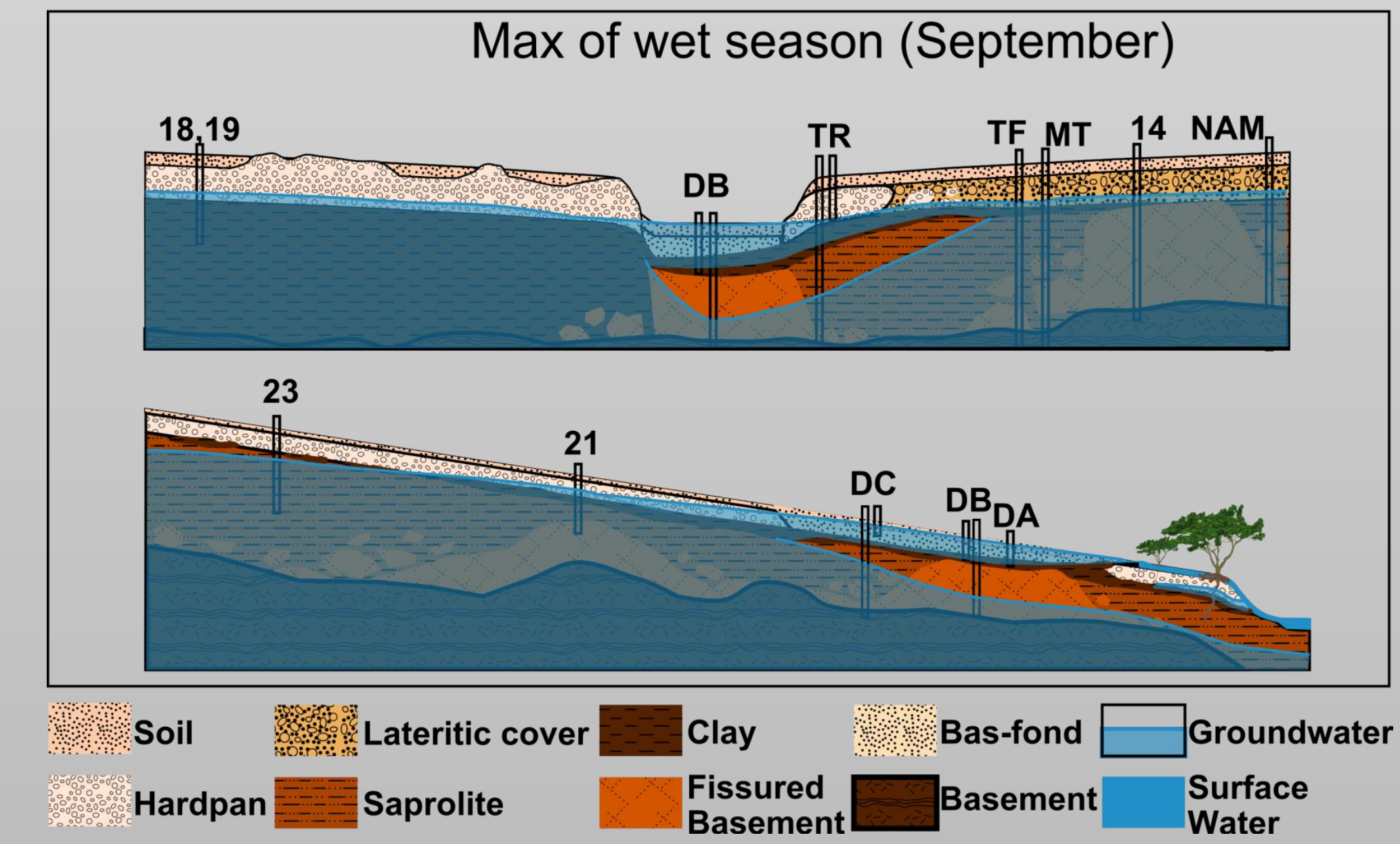
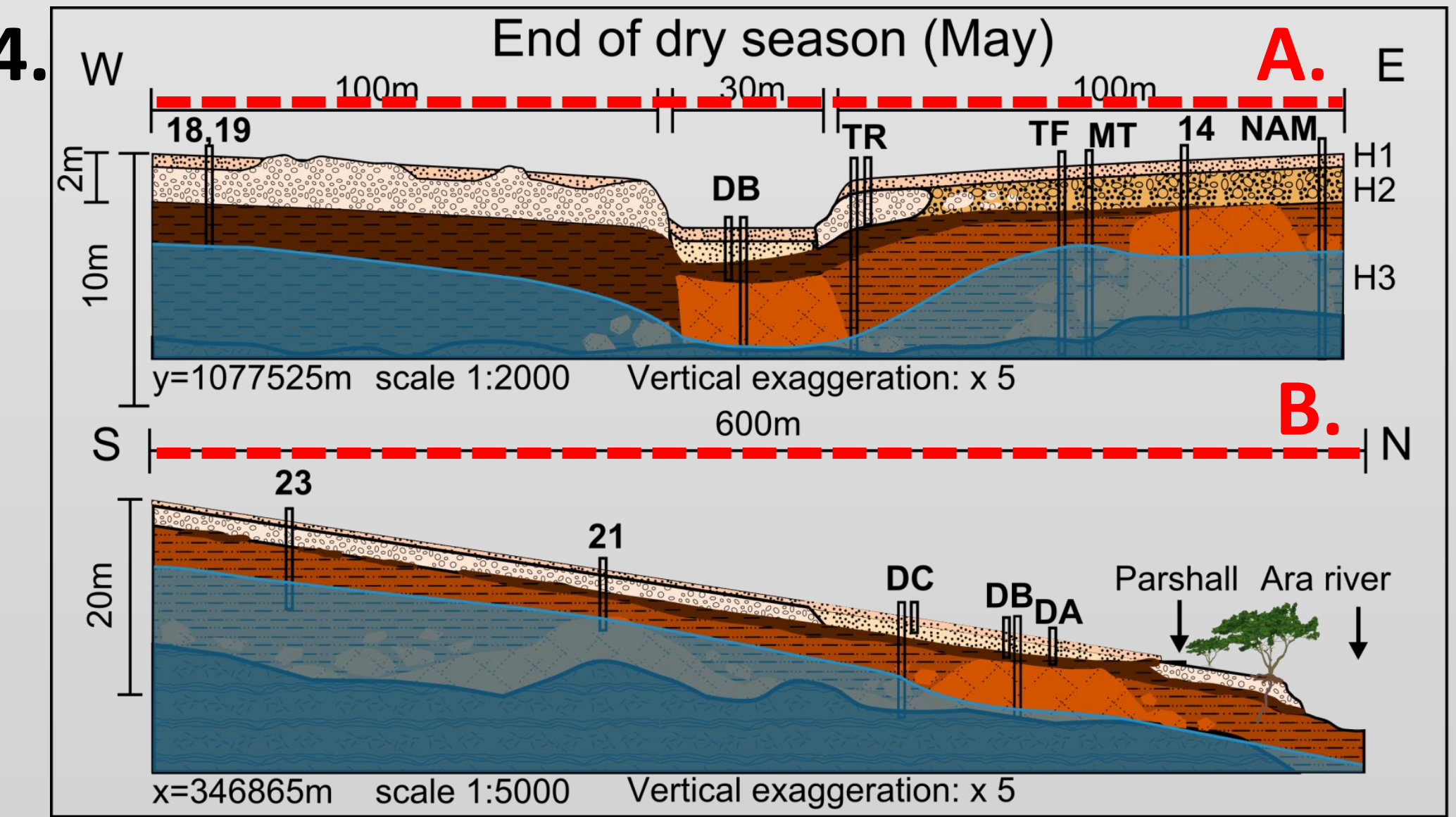


Water budget (ET = Evapotranspiration, Q = streamflow, S = Storage) for each year in mm and % of yearly Precipitation

Streamflow as semilog plot for the three tested cases

B. Conceptual model

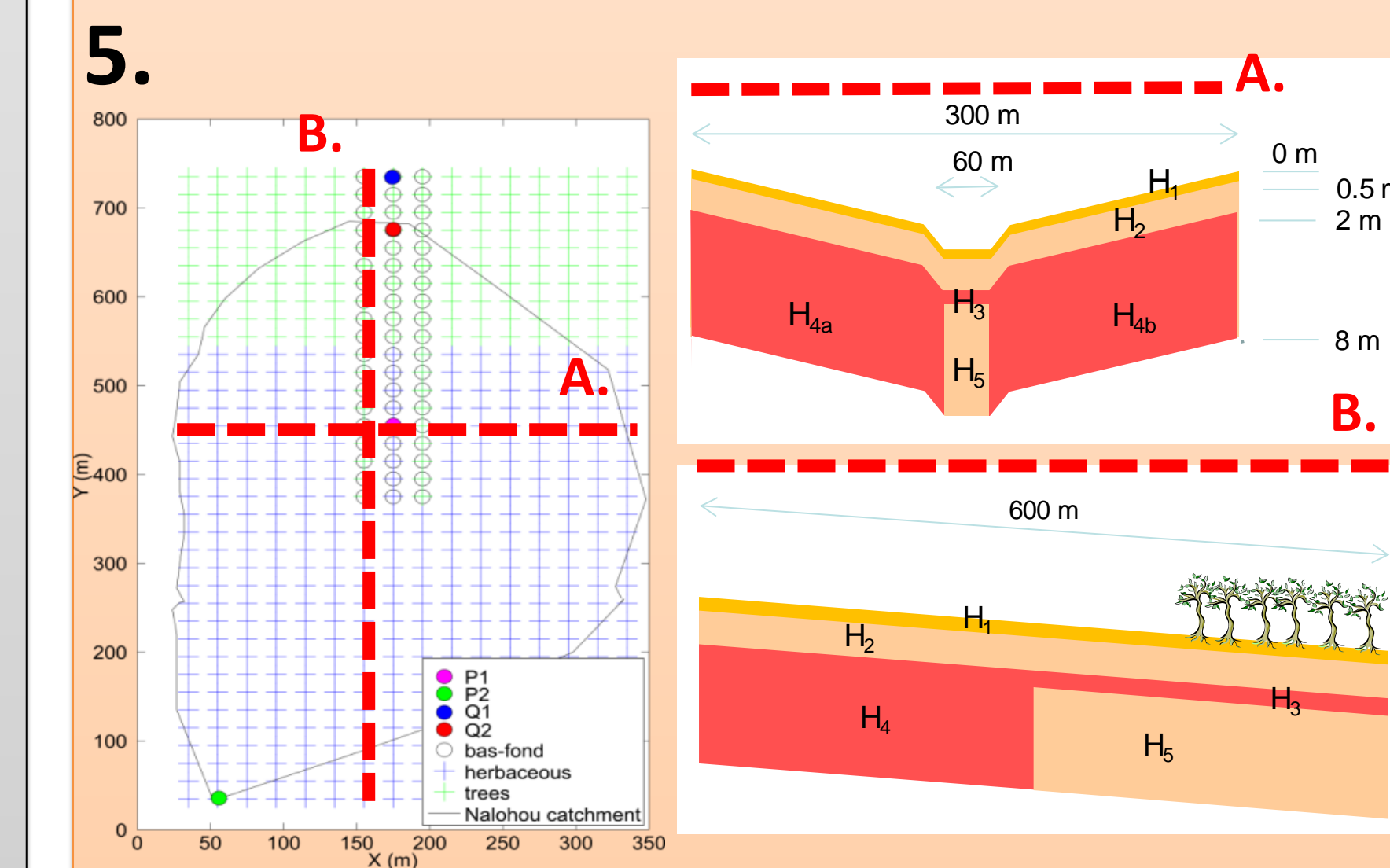
The conceptual model is built based on all observations. It is composed by a **fast system which feeds the baseflow through lateral connections** of perched water table and permanent water table in upstream areas; and a **slower system characterized by the seasonal oscillation of the permanent groundwater**. Permanent groundwater may not be connected to the river in the downstream areas, and is expected to discharge mostly into tree-driven evapotranspiration.



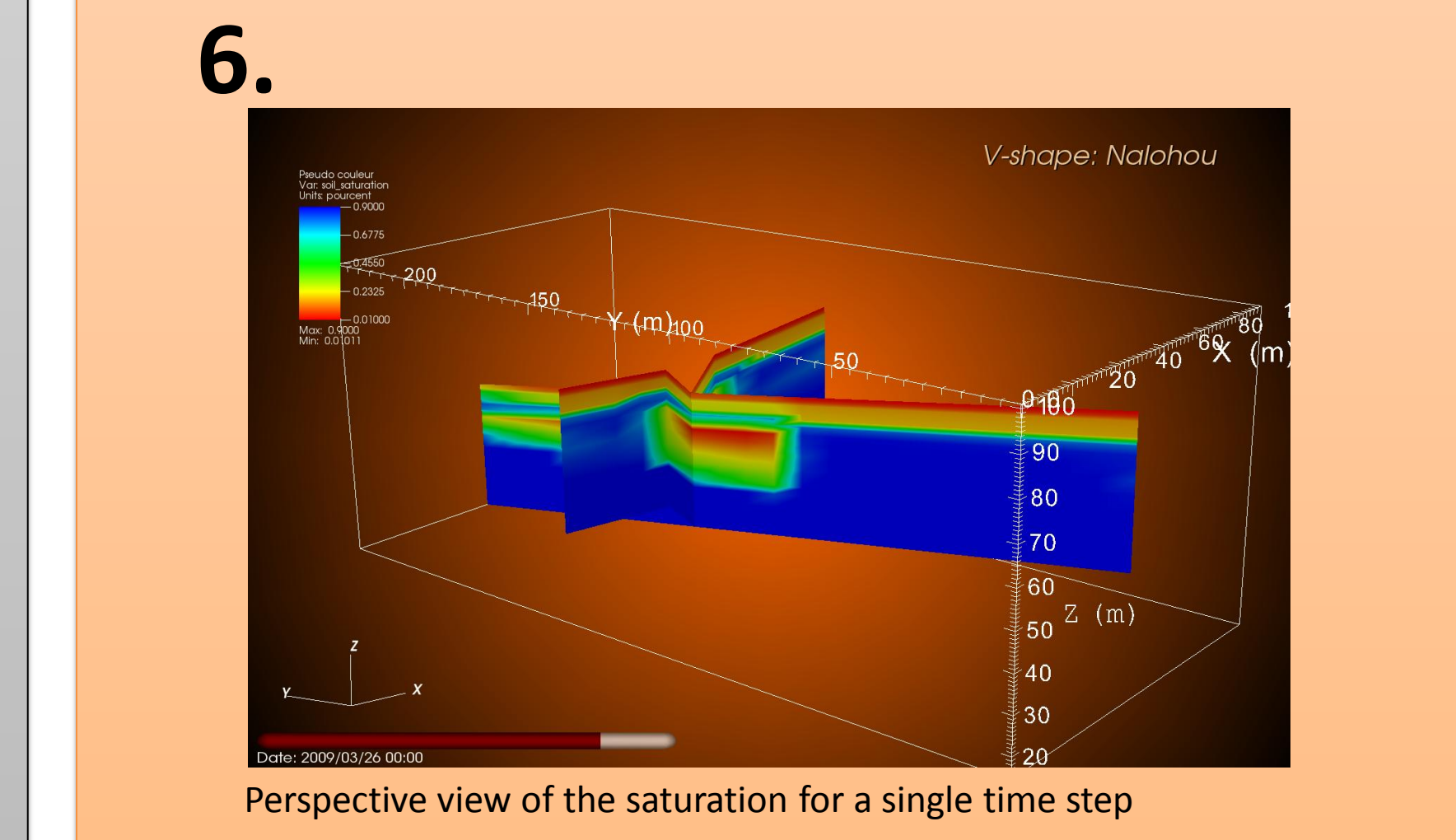
C. Critical zone model

Parflow is a saturated/unsaturated flow model with an overland flow scheme. It is coupled to a SVAT model: CLM, and allows to reproduce complex critical zone interactions within the terrestrial hydrological cycle for large scales and high resolutions (Maxwell & Miller, 2005).

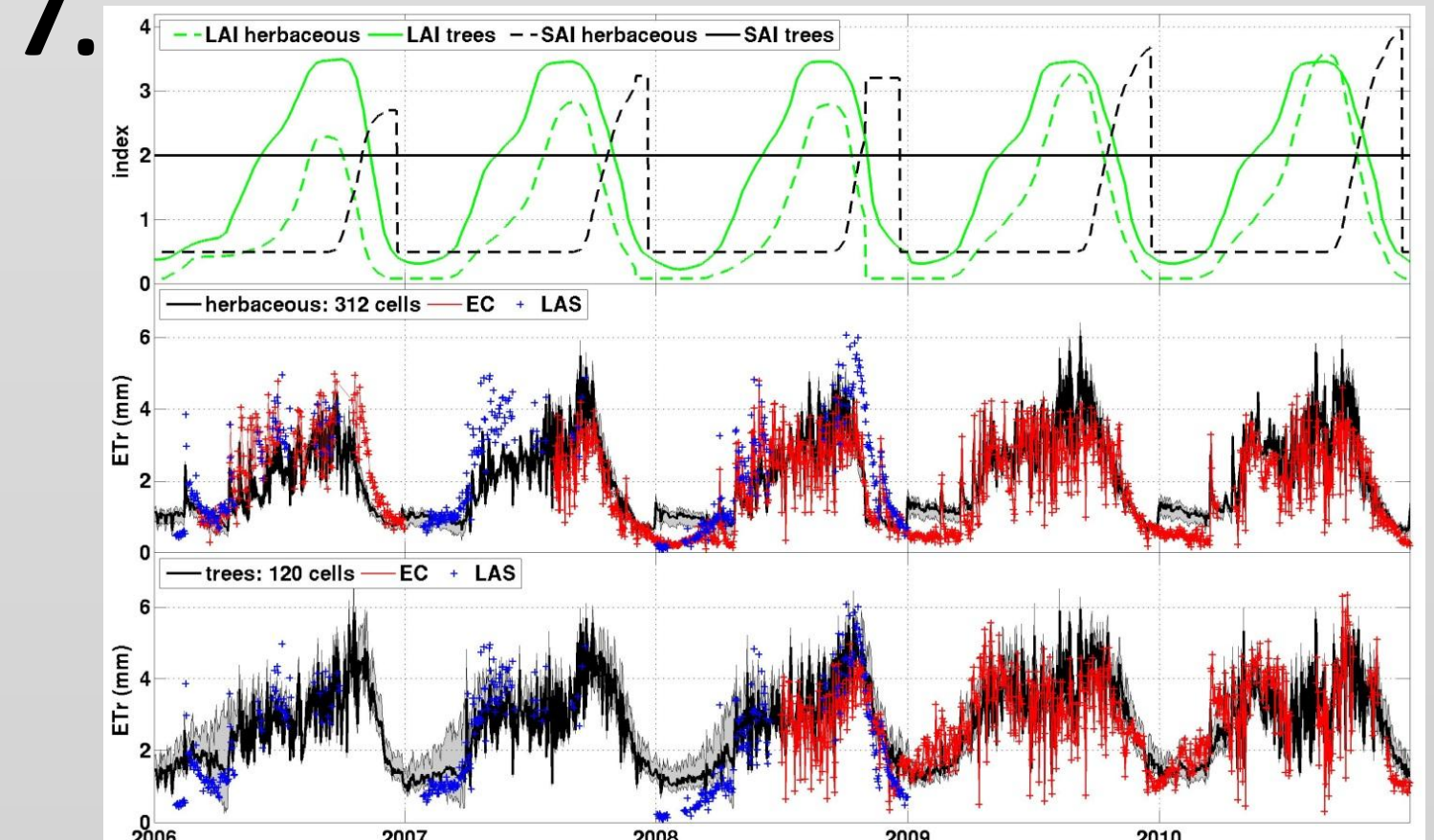
It is applied to a simplified (5) lithological model of the Nalohou catchment (1), from 2005 to 2012 with 30mn/20m resolution, forced with observed meteorological variables.



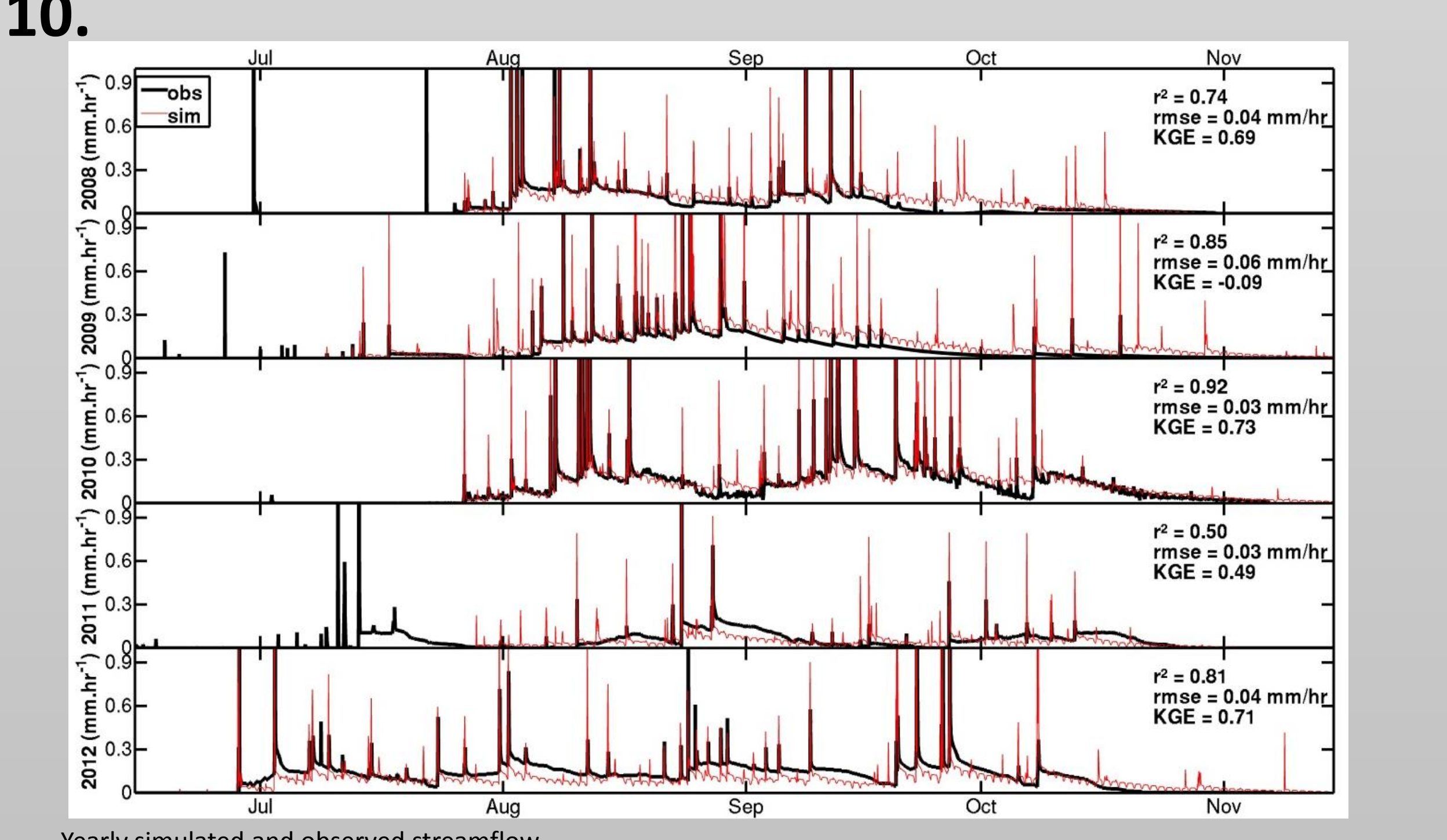
Left: catchment mesh. Blue cross = herb. Green cross = trees. Circles = Bas-fond. Right: Cross sections and lithological/pedological units



D. Model validation



Vegetation forcings (upper) and simulated and observed ET over herbaceous (center) and tree (lower panel) covers

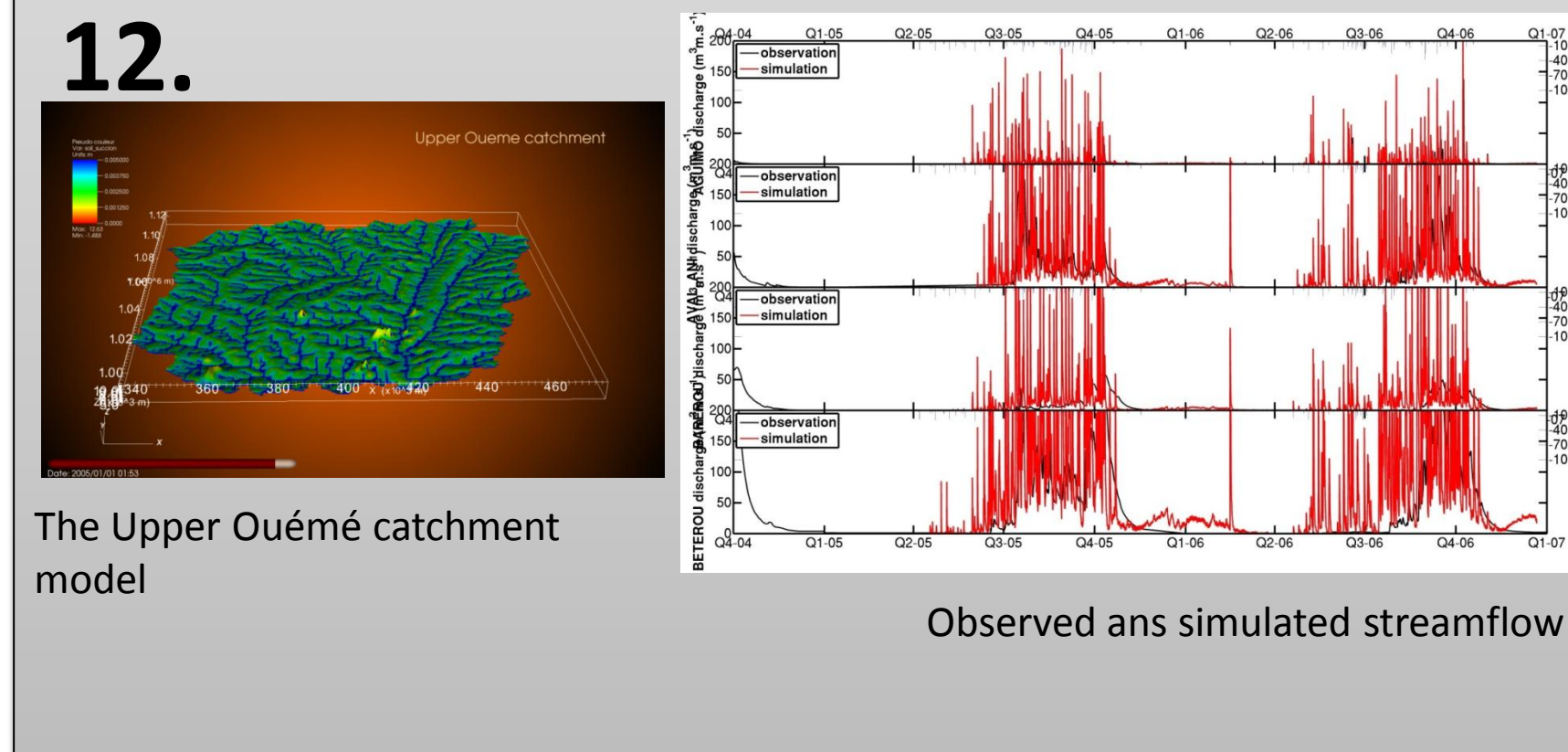


Yearly simulated and observed streamflow

- Good match for ET, except for too high simulated evaporation in the dry season (7). ET for trees is higher in the dry season and early wet season.
- The bas-fond shows the highest/left bank shows the lowest seasonal storage amplitude, both in simulation and gravity observation (8)
- Good match for streamflow onset and baseflow amplitude (10). High flow peaks do not affect much the budget because baseflow dominates largely (around 80%)

F. Meso-scale simulation

- The model is applied on the Upper-Oueme catchment (14300km²) with a 1km resolution (12)
- Large discrepancies with observation exist (13), but remain a satisfactory first step for a non-calibrated experiment



References

Maxwell, R.M., and Miller, N.L., 2005, Development of a Coupled Land Surface and Groundwater Model: Journal of Hydrometeorology, v. 6, p. 233–247, doi: 10.1175/JHM422.1.
 Hector, B., Séguis, L., Hinderer, J., Cohard, J.-M., Wubda, M., Desclotres, M., Benarros, N., and Boy, J.-P., 2015, Water storage changes as a marker for base flow generation processes in a tropical humid basement catchment (Benin): Insights from hybrid gravimetry: Water Resources Research, doi: 10.1002/2014WR015773.
 Hector, B., Cohard, J.-M., Maxwell, R.-M., Séguis, L., Galle, S.: Hydrological functioning of West-African inland valleys explored with a critical zone model. Hydrol. Earth Syst. Sci. In prep.

Conclusions

- A physically-based critical zone model is able to simulate states and fluxes satisfyingly even in highly heterogeneous hard-rock basement with intermittent hydrology.
- Such explicit representation of the critical zone allows to conduct virtual experiments.
- In the Sudanian context, replacing trees by human-induced herbaceous cover may decrease the evapotranspiration of as much as 13% of yearly precipitation and increase streamflow significantly.
- The major streamflow decrease since the 80' is due to the precipitation decrease, but the impact of the concomitant deforestation is not clear yet.
- The impact of such global changes at meso-scale basins will be investigated.