

Importance of immobile water for transport in porous aquifers

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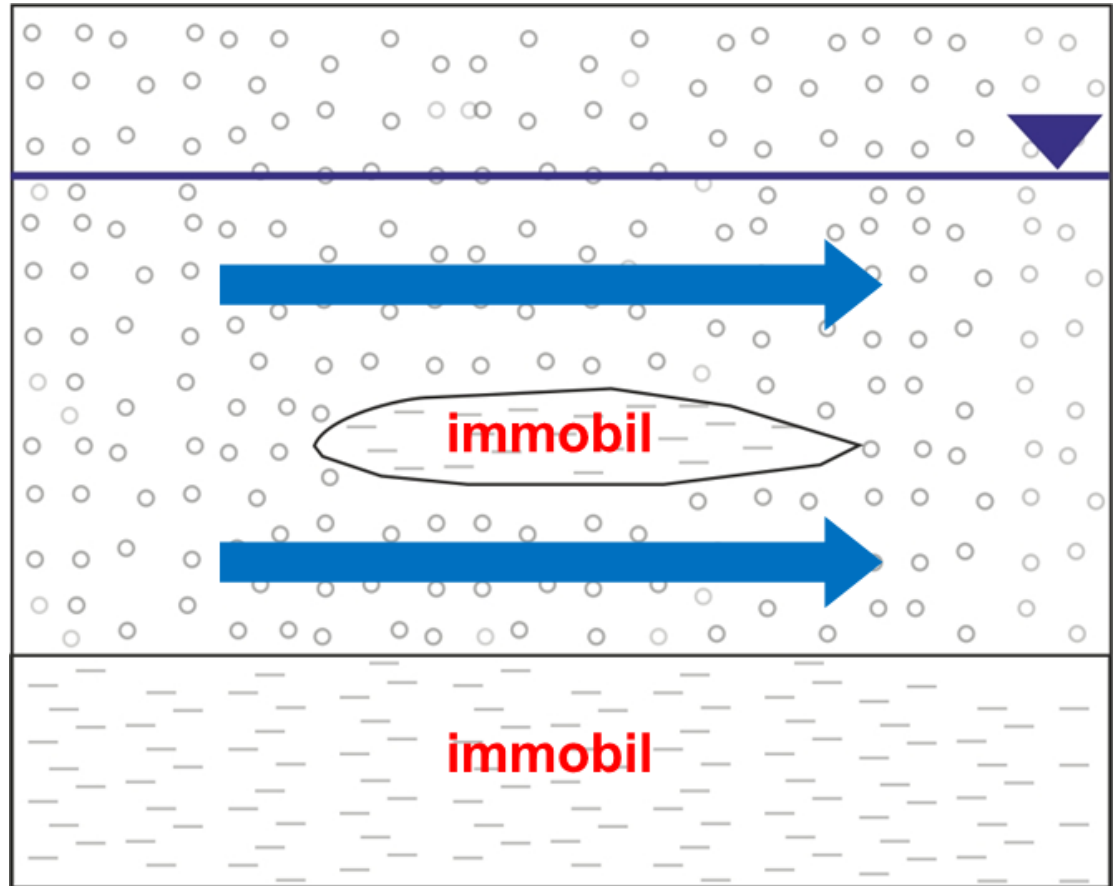
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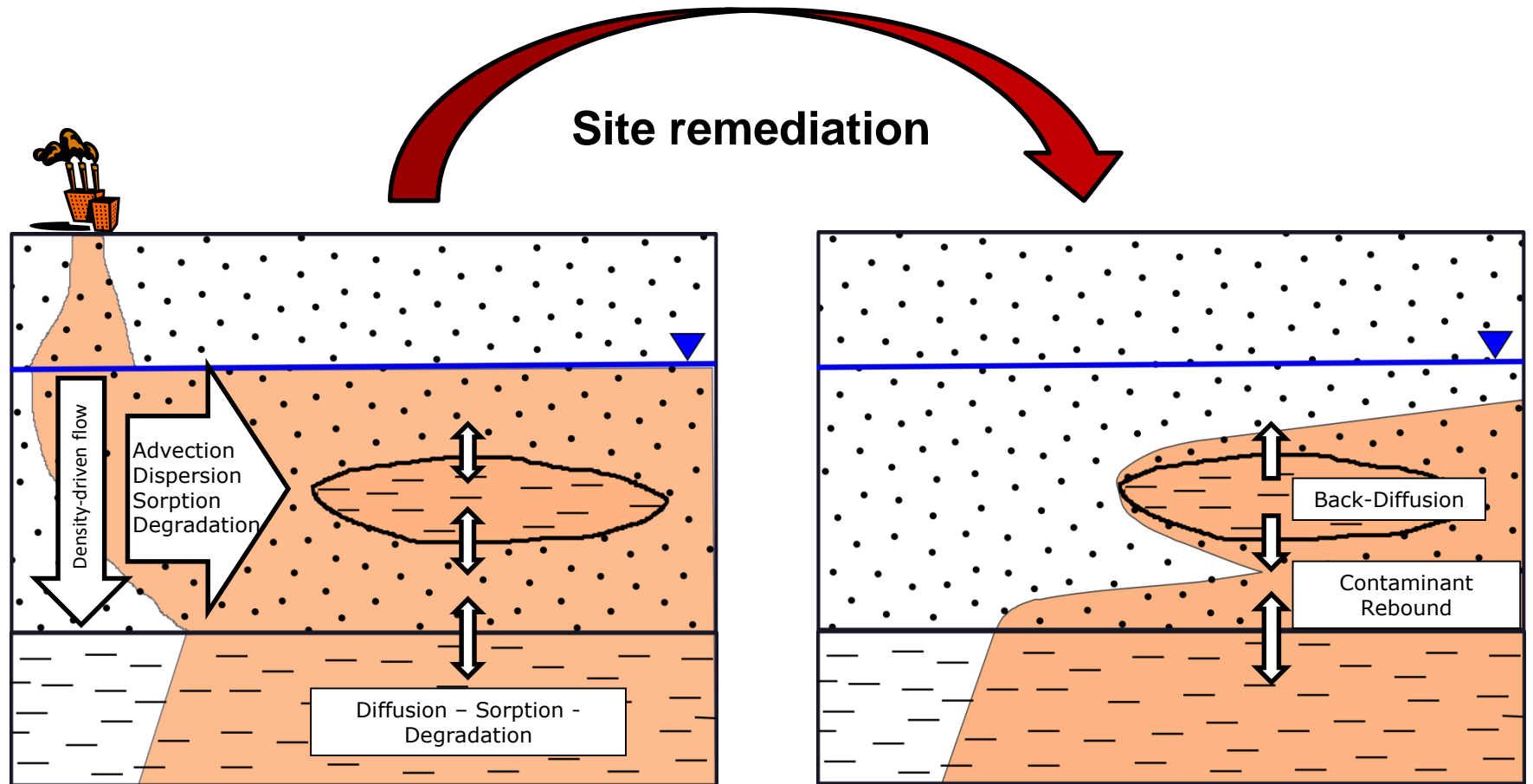
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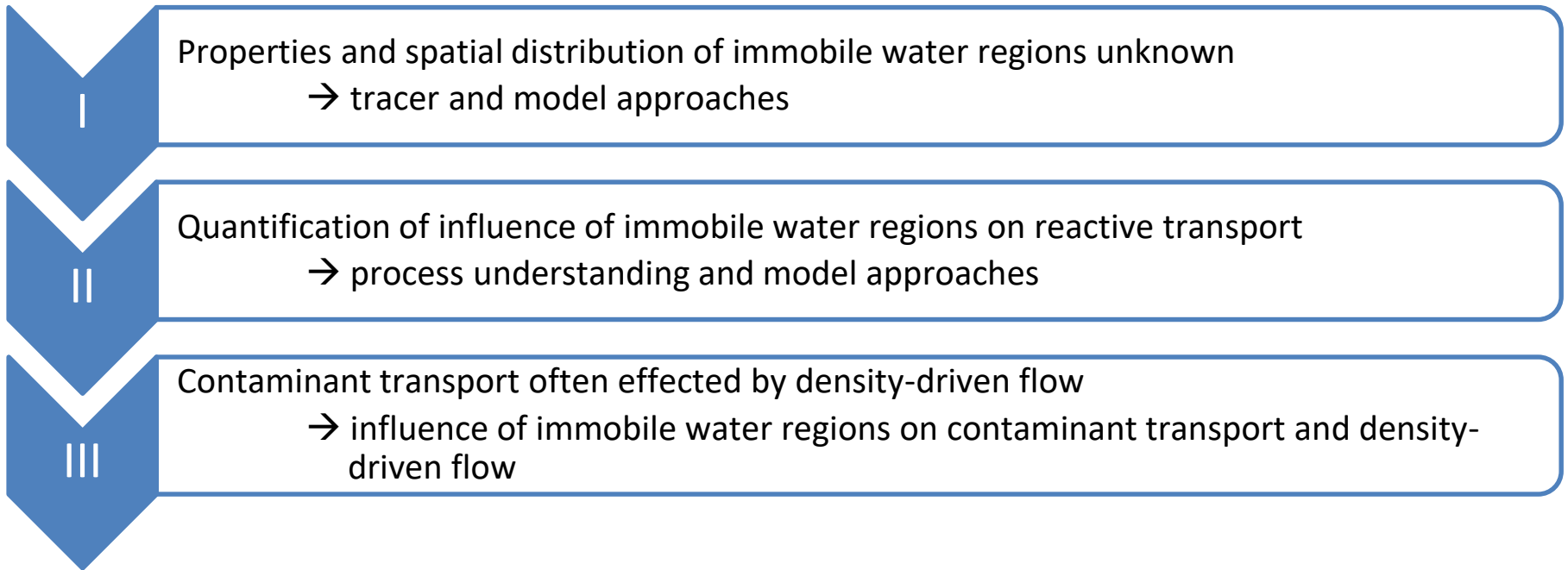
Aquifers are heterogeneous containing less permeable layers



Impact of immobile water on transport



Research gaps

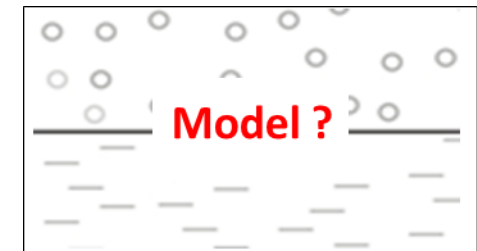
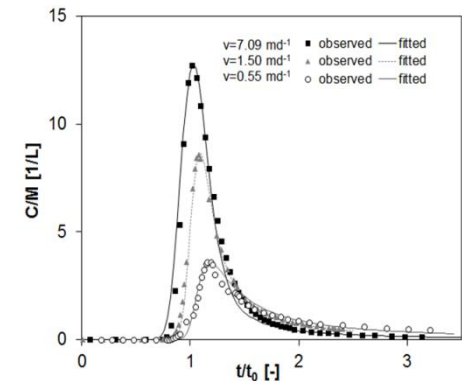
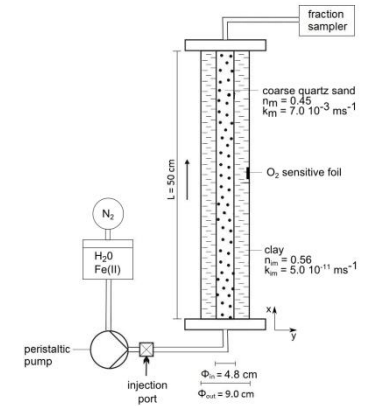


Objectives

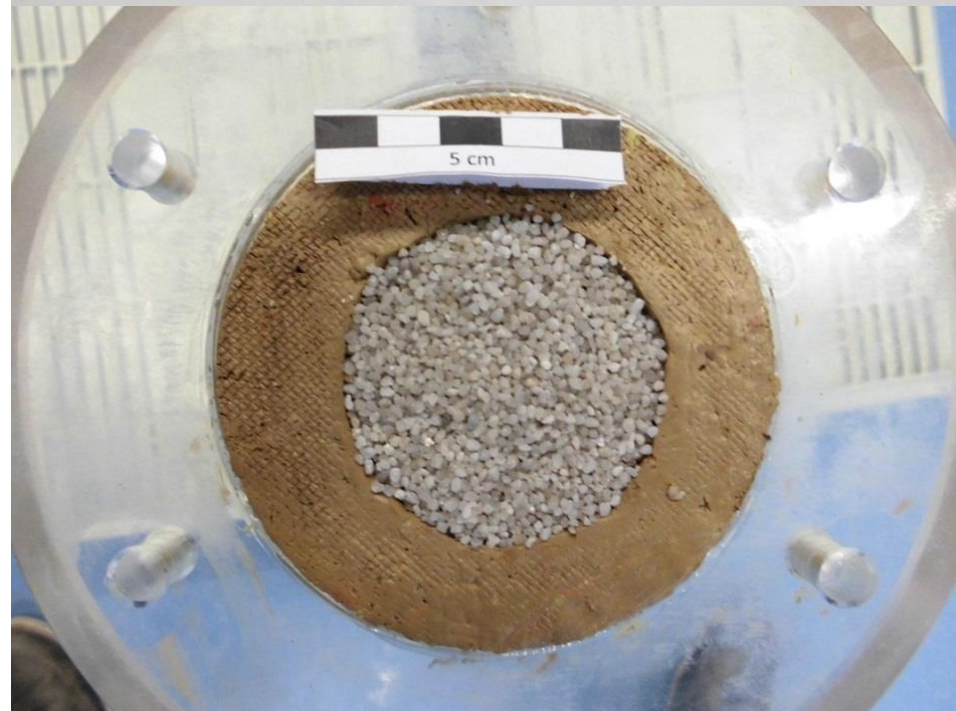
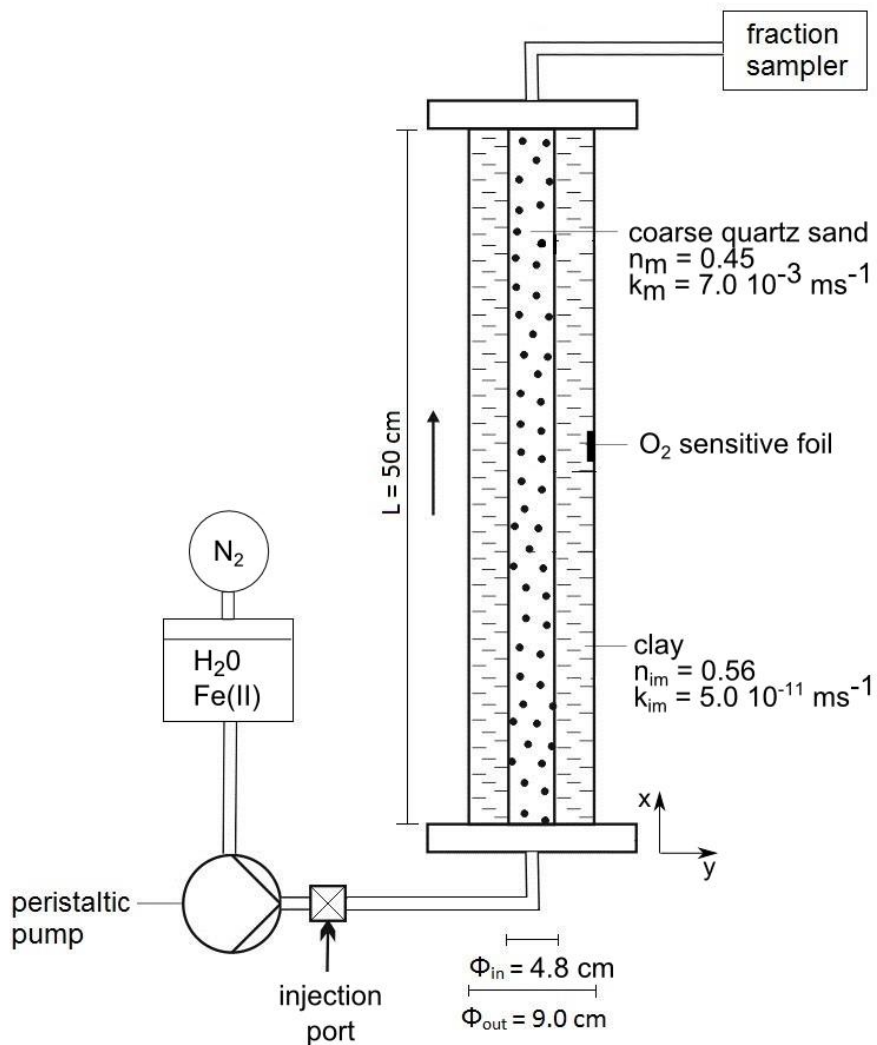
1) develop experimental methods to study specific transport processes in dual-porosity systems

2) identify impact of immobile water on transport processes in dual-porosity systems

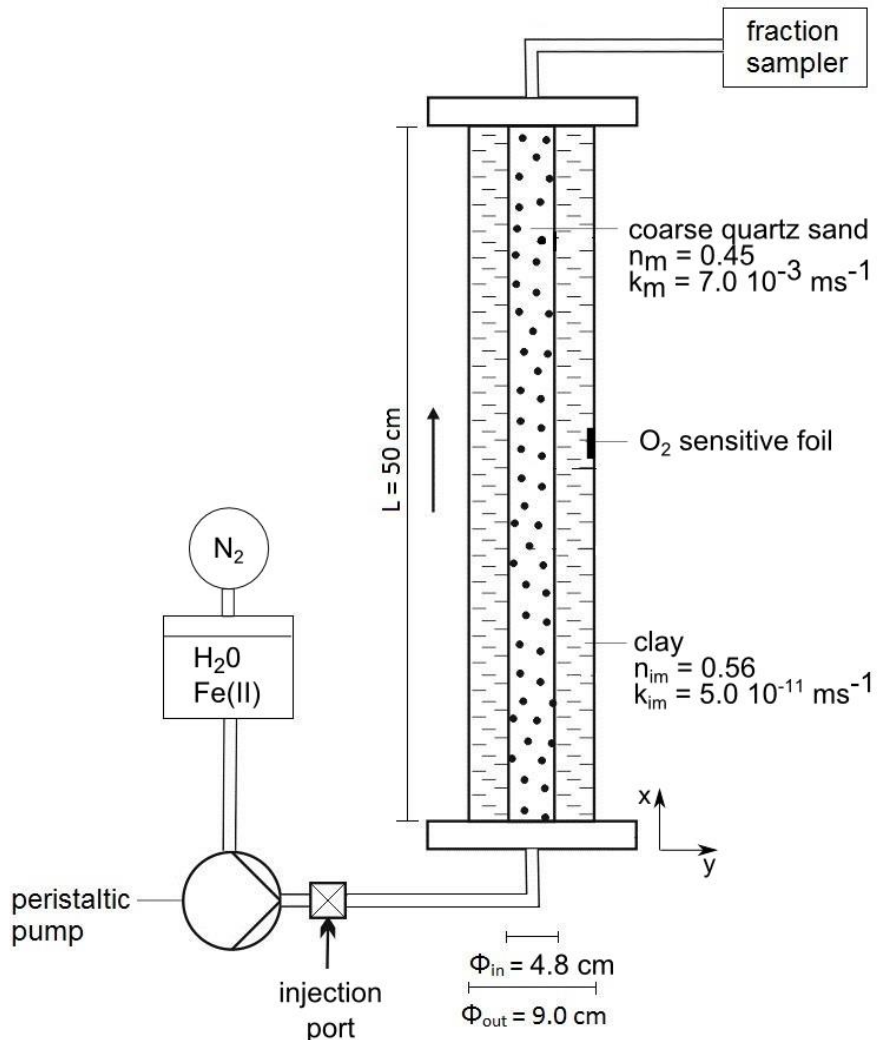
3) improve mathematical models for reactive transport modelling in dual-porosity systems



Experimental Setup



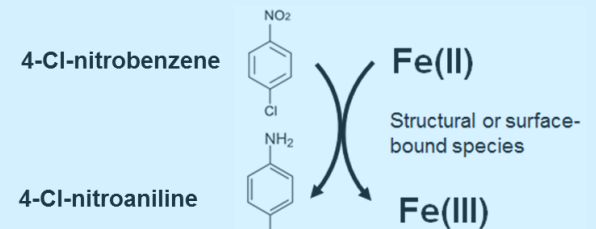
Experimental Setup



1) conservative transport

$$D_{\text{uranine}} < D_{\text{bromide}} < D_{\text{deuterium}}$$

2a) reactive transport



Source: Hofstetter et al. (1999), Elsner et al. (2004)

2b) reactive transport

$^{134}\text{Cesium}$

3) density driven flow

Methanol

$$0 > -0.0036 > -0.0071$$

Mathematical Modelling

Single Fissure Dispersion Model (SFDM) including sorption and degradation/decay:

$$\frac{\partial C_m}{\partial t} + v \frac{\partial C_m}{\partial x} - D \frac{\partial^2 C_m}{\partial x^2} - \frac{n_{im} D_{im}}{2b} \frac{\partial C_{im}}{\partial y} \Big|_{y=\frac{\phi_{in}}{2}} = 0 \quad \text{for} \quad 0 \leq y < \frac{\phi_{in}}{2}$$

$$\frac{\partial C_{im}}{\partial t} + R_{im} \frac{\partial C_{im}}{\partial t} - D_{im} \frac{\partial^2 C_{im}}{\partial y^2} + k C_{im} = 0 \quad \text{for} \quad \frac{\phi_{in}}{2} \leq y < \infty$$

Analytical Solution

Source: Maloszewski and Zuber (1985, 1990)

$$C_m(t) = \frac{M a \sqrt{R_{im}}}{2\pi Q} \sqrt{\frac{t_0}{P_D}} \int_0^t \exp \left[-\frac{(t_0 - u)^2}{4P_D u t_0} - \frac{(a\sqrt{R_{im}})^2 u^2}{t - u} - \frac{k_{degr}}{R_{im}} (t - u) \right] \frac{du}{\sqrt{u(t - u)^3}}$$

Fitting Parameter:

transit time of water

$$t_0 = \frac{V_m}{Q}$$

diffusion parameter

$$a = \frac{n_p \sqrt{D_p}}{2b}$$

dispersion parameter

$$P_D = \frac{\alpha_L}{L}$$

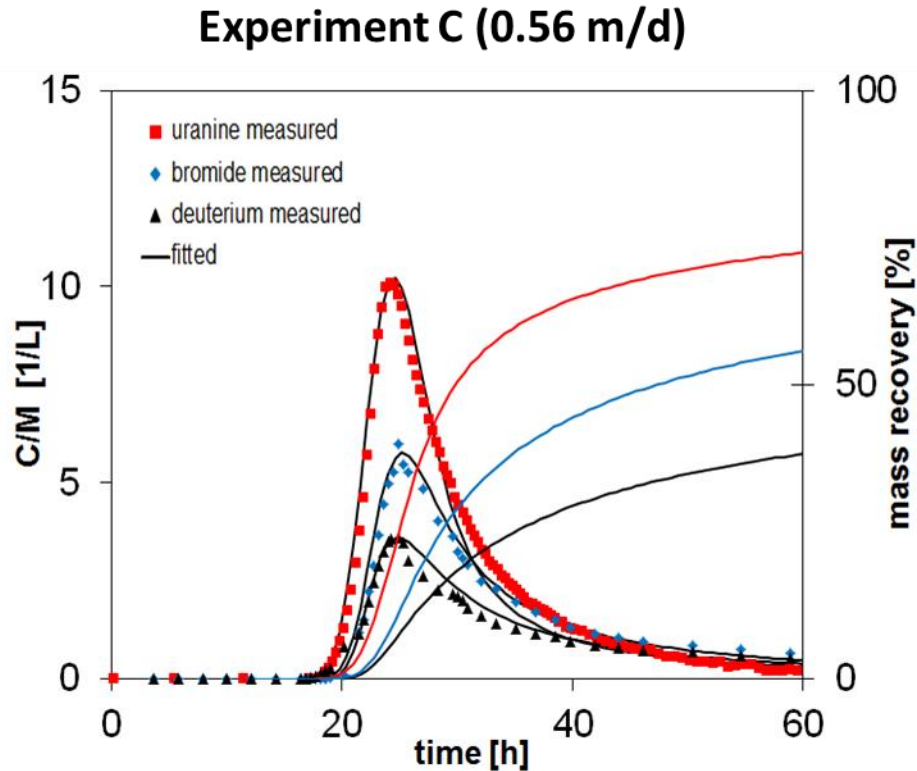
retardation factor

$$R_{im} = 1 + \frac{\rho_b}{n_p} K_d$$

first-order rate

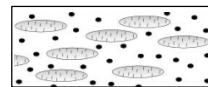
$$k_{degr}$$

Results – conservative transport

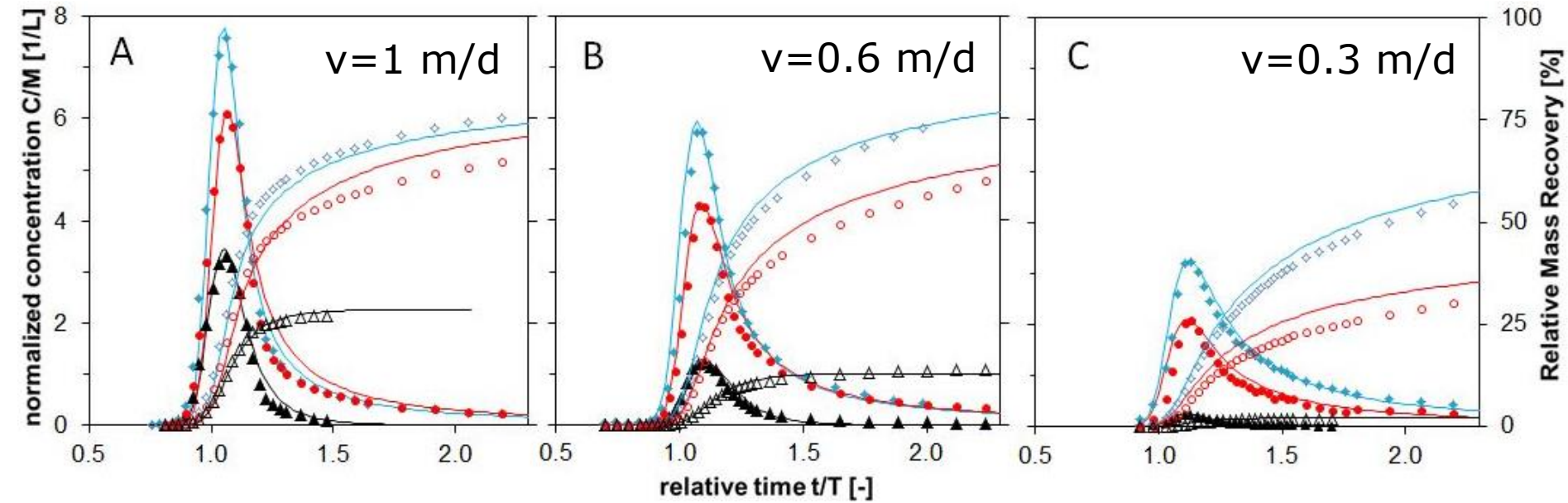


(Knorr et al. Hydrological Processes, 2016)

- fit parameter in agreement with measurements ($n_{\text{mob}} + n_{\text{immob}}$)
- also applicable for randomly distributed immobile water regions



Results – reactive transport: 4Cl-Nitrobenzene



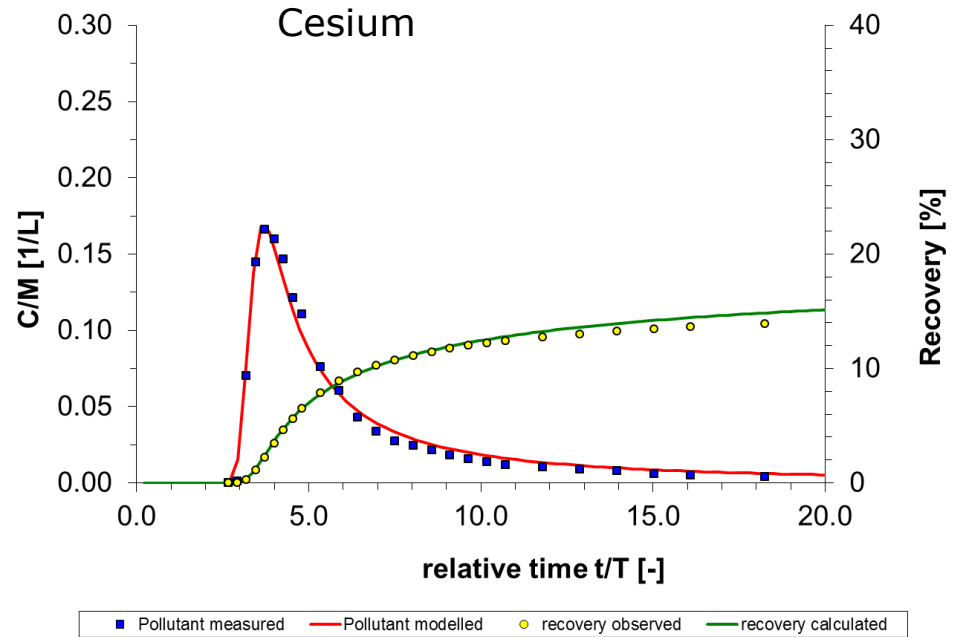
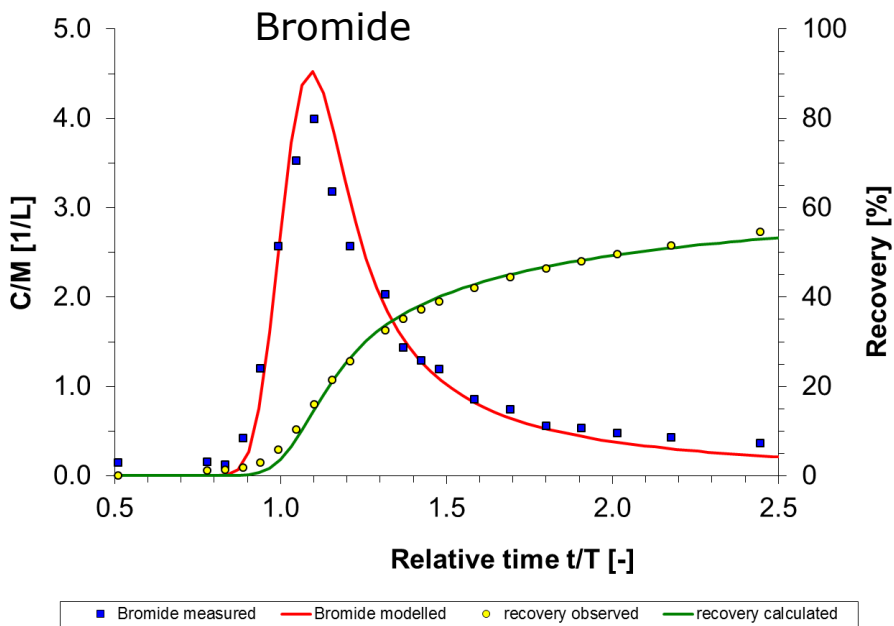
$$RR_{(4\text{-Cl-AN})} = 12\%$$

$$RR_{(4\text{-Cl-AN})} = 17\%$$

$$RR_{(4\text{-Cl-AN})} = 28\%$$

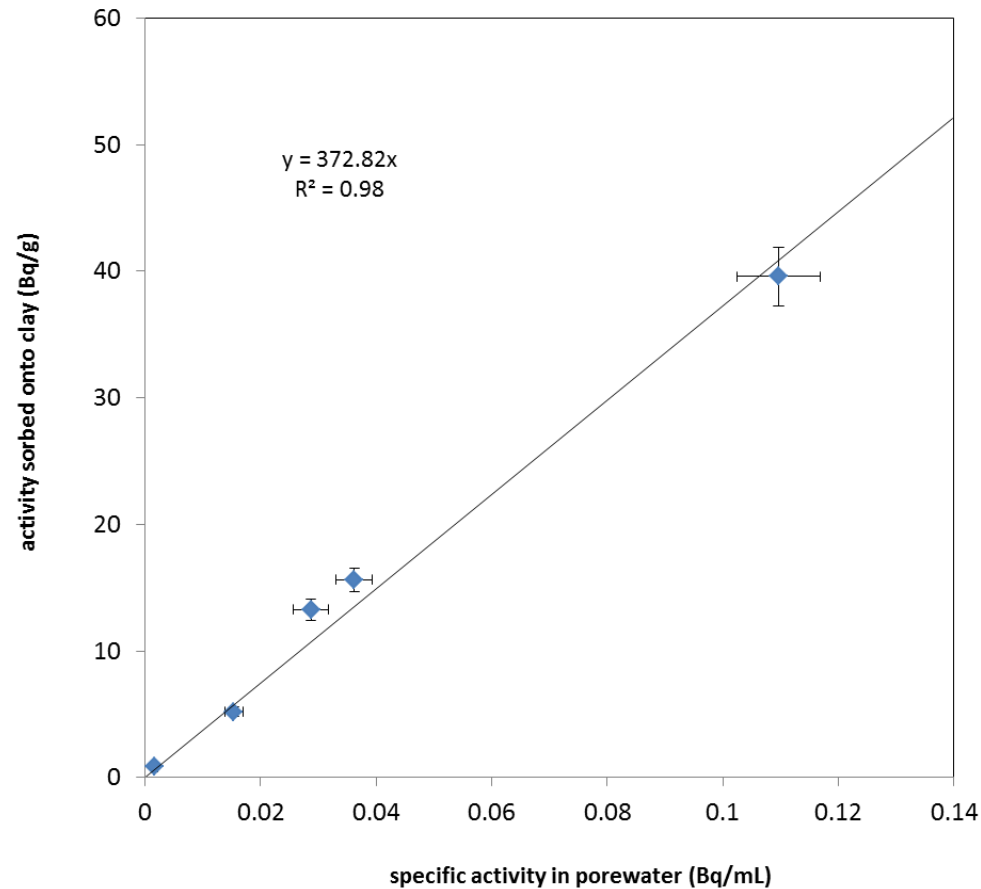
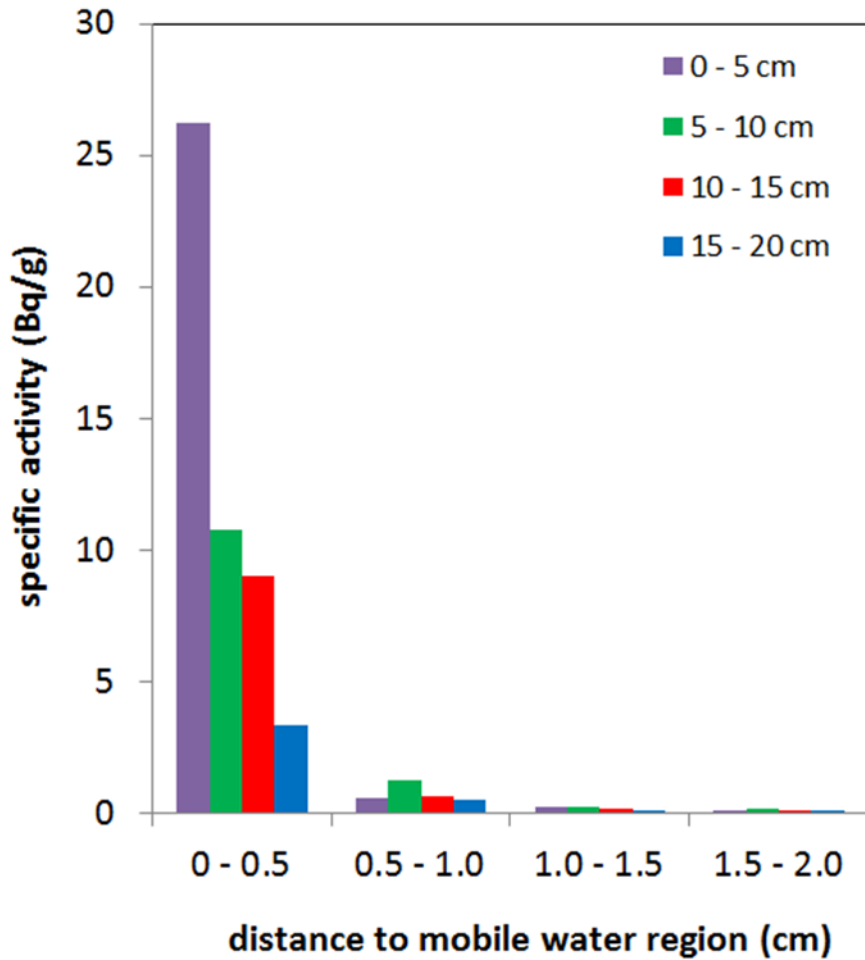
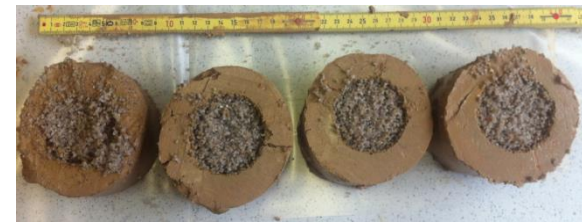
- fit parameter in agreement with measurements ($n_{\text{mob}} + n_{\text{immob}}$)
- back diffusion limited by transformation
- transformation independent on flow velocity but on transit time

Results – reactive transport: ^{134}Cs



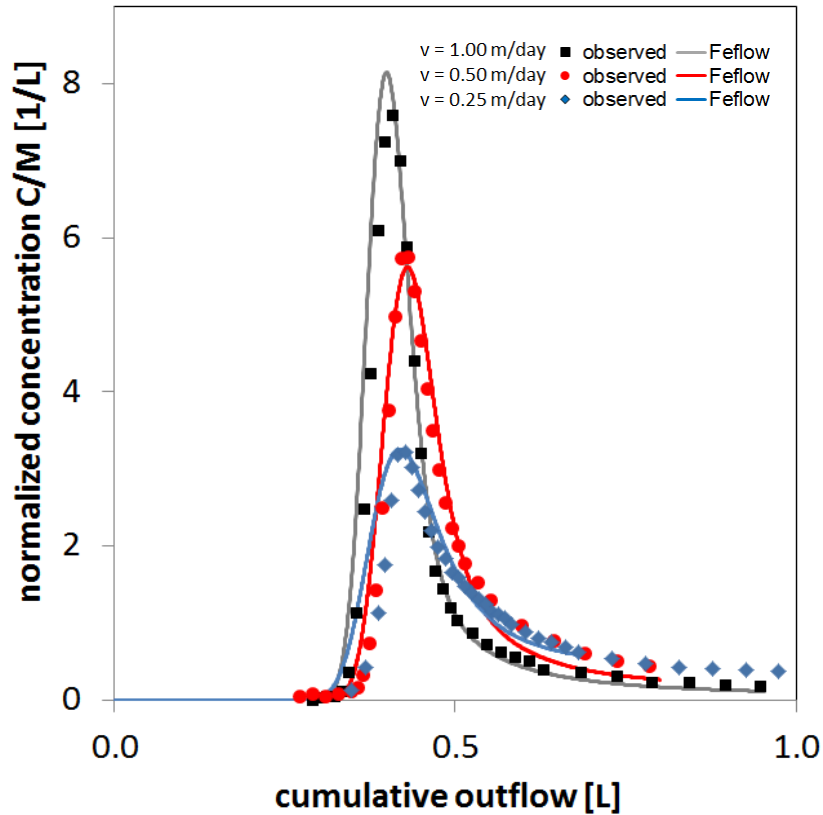
- strong retardation of ^{134}Cs ($R=3.2$)
- impact on diffusion into immobile water
- same immobile water content for bromide and cesium

Results – reactive transport: ^{134}Cs

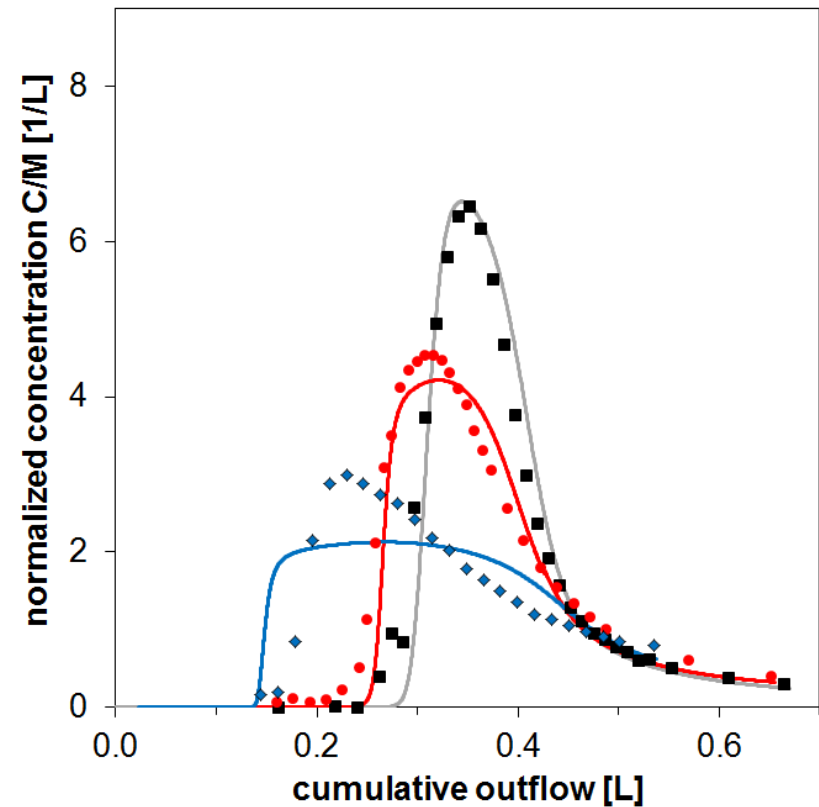


Results – density driven flow

Model Calibration



Density-driven flow (4% Methanol $\Delta\rho = -0.0071$)

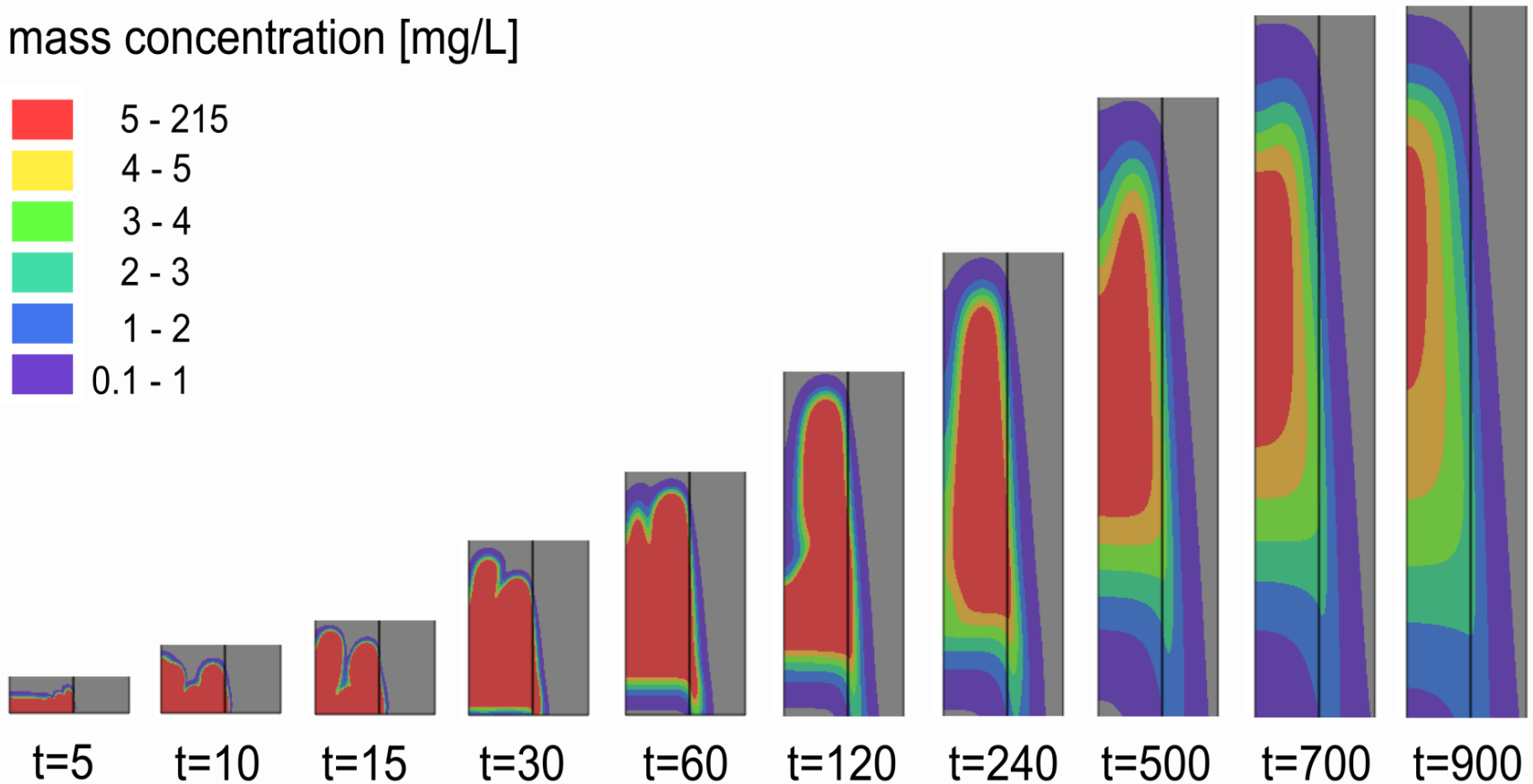


- very small differences in density caused instabilities
- increased diffusion due to longer contact times
- numerical modelling required

Results – density driven flow Feflow simulation

2D axisymmetric

mass concentration [mg/L]



Conclusions

- immobile water region strongly impact transport in heterogeneous aquifers
- mass transfer depends on diffusion coefficient
- multitracer approach
- abiotic reaction reduces mass transfer
- mass transfer enhanced for density driven flow and inhibition of flow instabilities
- SFDM also applicable in porous aquifers for conservative and reactive transport
- numerical modelling required to describe density driven flow