

Permeability in deep-seated granitic rocks: lessons learnt from deep geothermal boreholes in the Upper Rhine Graben (URG)



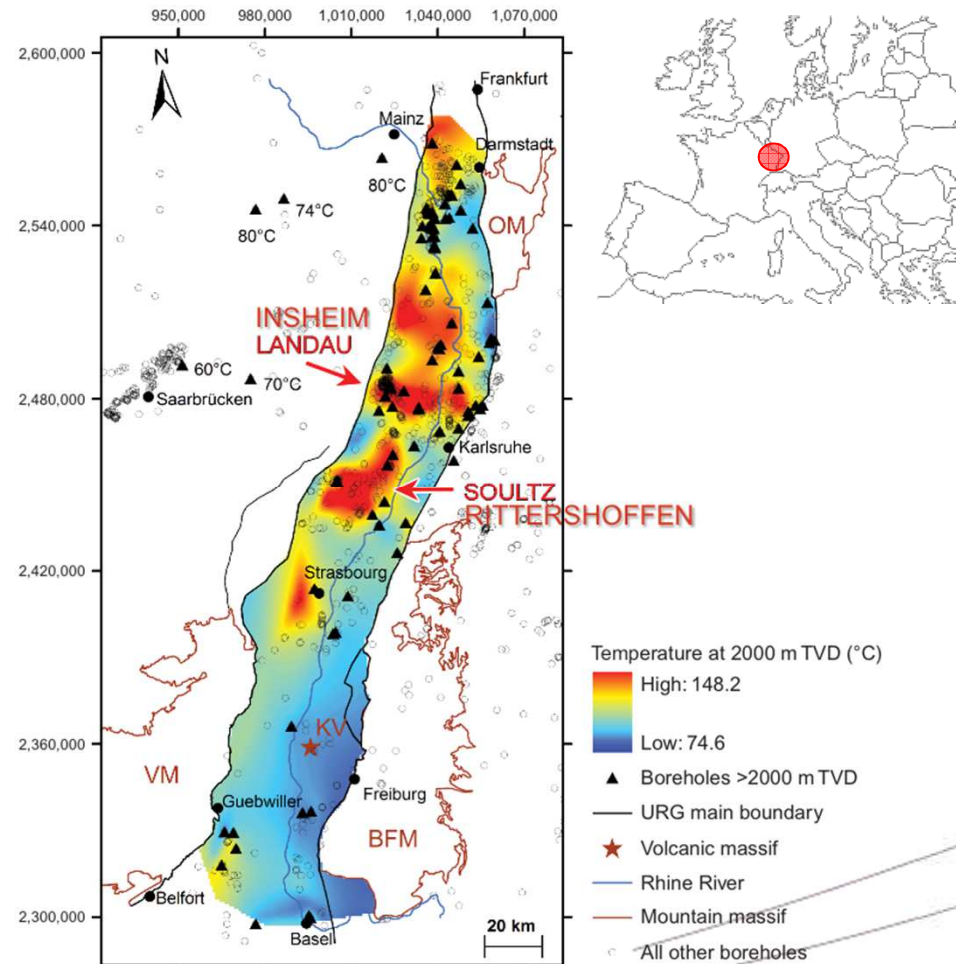
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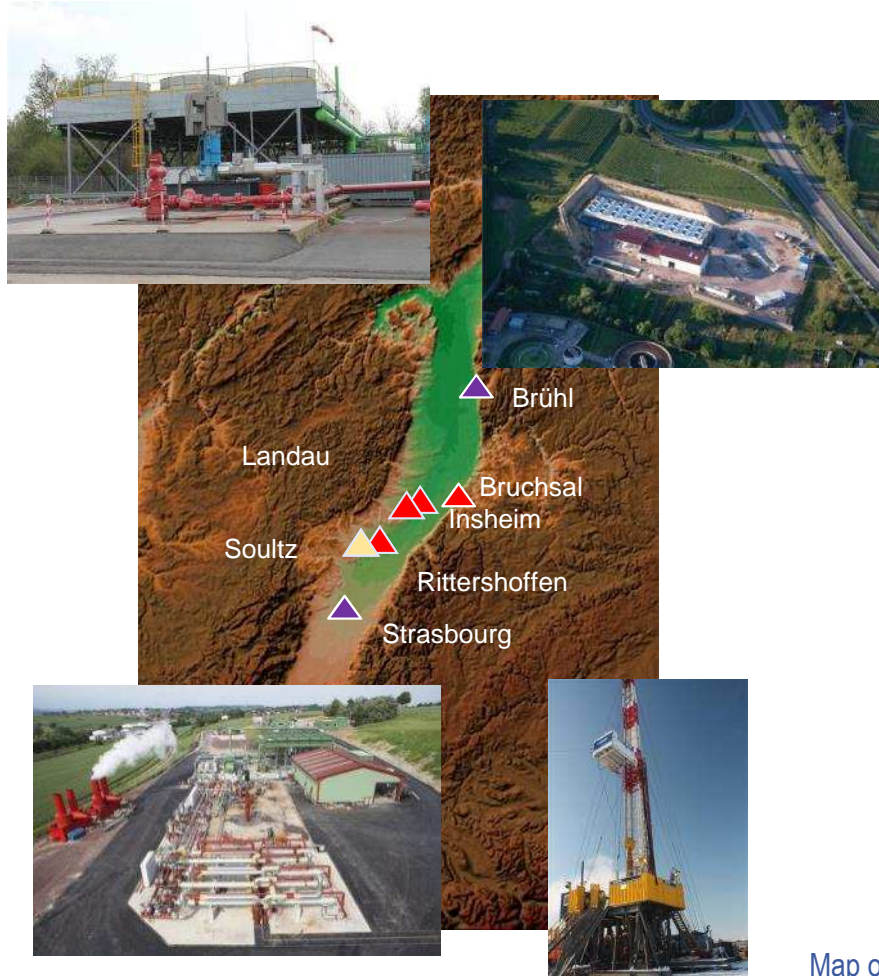
Context of the study

- Several thermal anomalies in Upper Rhine Graben (Europe) associated with **hydrothermal circulations in the Triassic clastic sediments and Paleozoic crystalline basement**
- Geothermal boreholes drilled from 2.5 to 5 km depth
→ Bottom hole temperatures range from **170°C to 200°C**
- **Water bearing fractures** characterized by geothermal brines (TDS 100g/L)



Temperature map in the URG (Bailleux et al., 2013)

Geothermal projects in the URG

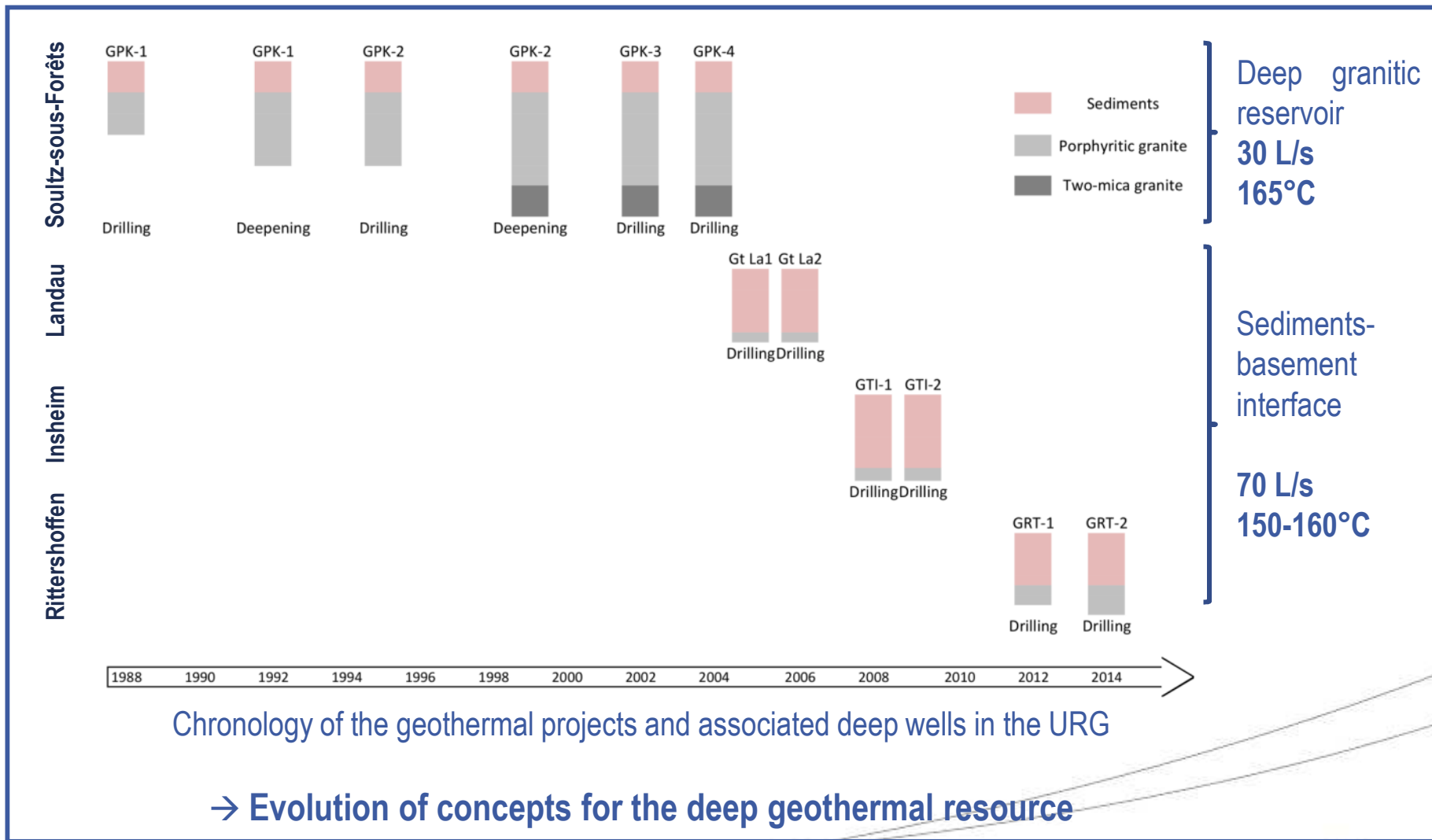


- **Experimental pilot project**
Sultz-sous-Forêts
- **Industrial projects** based on lessons learned from Sultz
Landau, Insheim, Rittershoffen
- **Several ongoing projects**
4 projects around Strasbourg

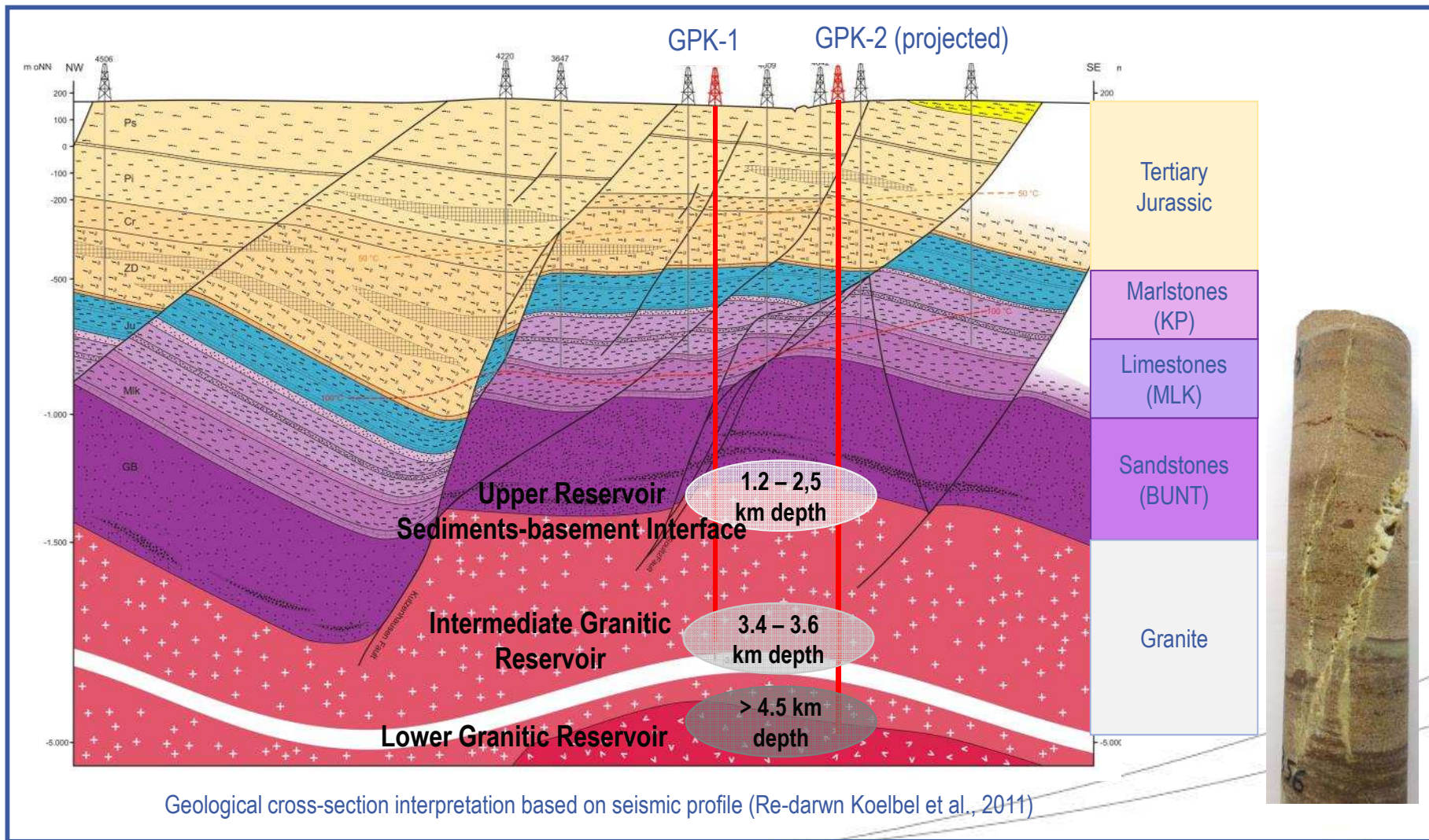
→ URG is the **most active zone** for deep geothermal energy exploration and exploitation in France

Map of the geothermal projects in the URG

Deep geothermal wells in the URG



Soultz geothermal project



Permeable natural fractures

- **EPS-1** = exploration well → 800 m of core samples + borehole images

More than **3000 natural fractures** observed in Triassic sediments and in the **granitic basement**

- All fractures show **hydrothermal fillings** (barite, galena, secondary quartz, clays...)

Fillings are **partial** and **channelized**

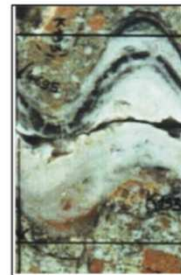
→ **Less than 1% of fracture are permeable** at the borehole scale

- Sealing is caused by **alteration of granite** and **circulation of brine**

CORE SAMPLE



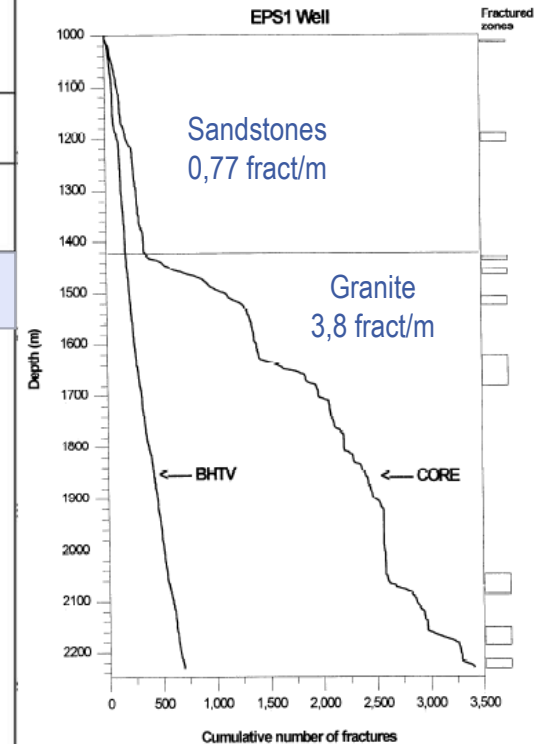
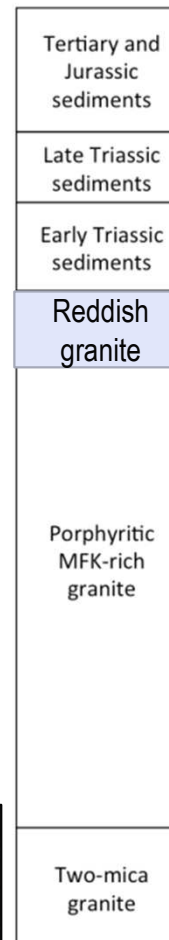
Fracture zone in the well EPS-1 at 1220 m depth, filled with barite and galena and correlating partial mud losses



Fracture zone in the well EPS-1 at 2156 m depth, filled with secondary quartz and correlating total mud losses and temperature variation



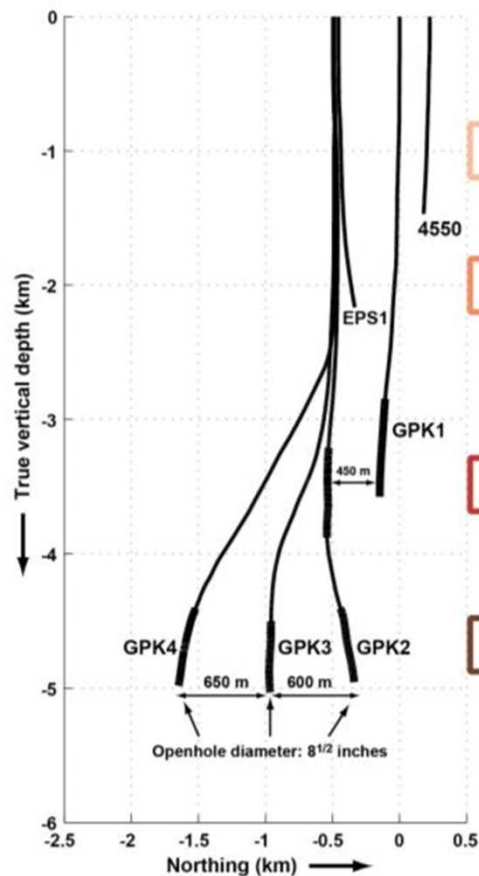
Matrix permeability = 10^{-18} - 10^{-21} m²



Fracture density:
Granite >> Sandstone

(after Genter et al., 1997)

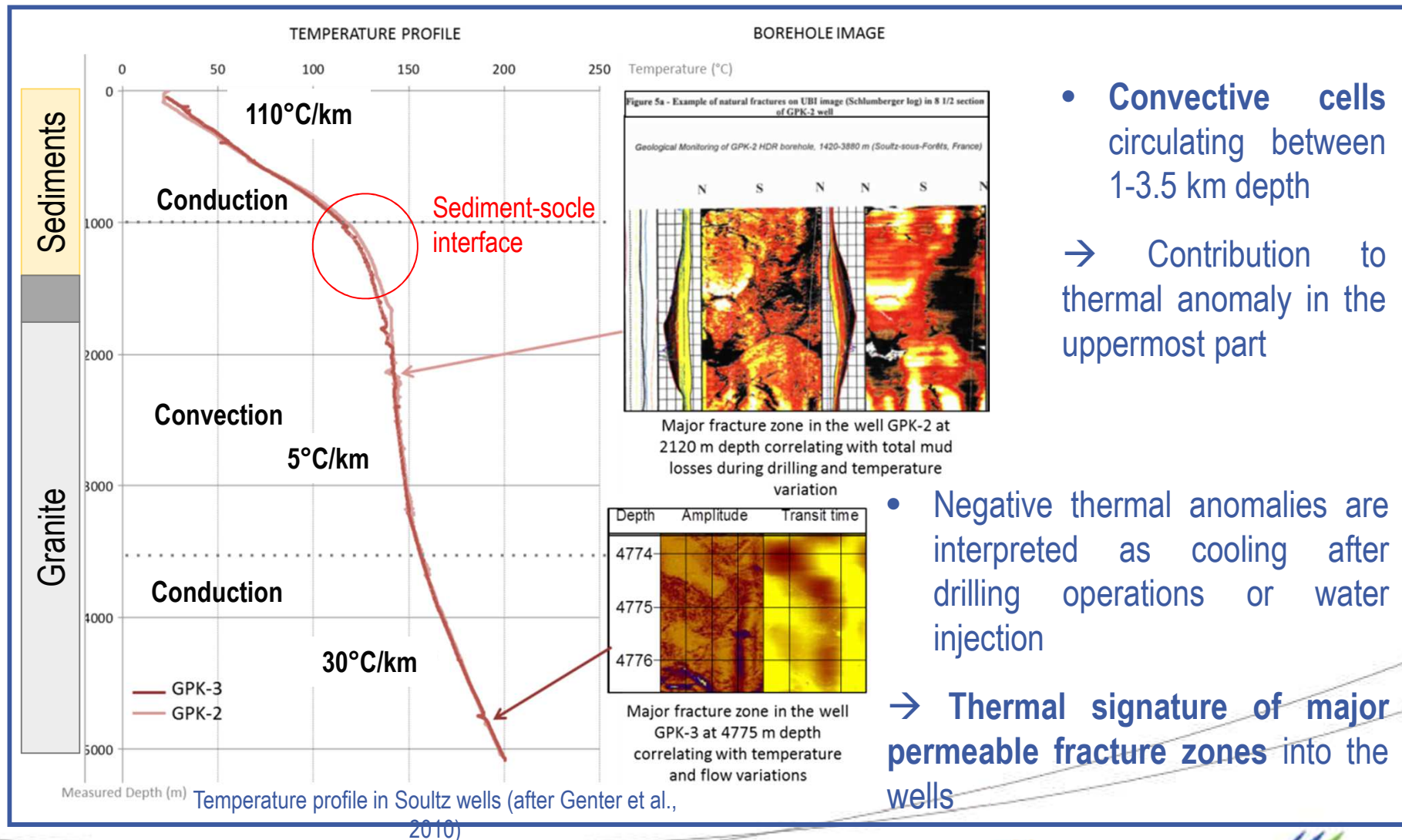
Natural injectivity of the Soultz wells



N-S cross-section through Soultz site

- Natural injectivity **very variable** in reservoirs and in Soultz wells
- 3 L/s/bar
- Major fracture zones with total mud losses during drilling operations
- Top granite = 200 m of reddish granite paleoemersion → paleoweathering process and fractured zones
- 0,01-0,02 L/s/bar
- Need of stimulation of deep granitic reservoir

Natural hydrothermal circulation



- Convective cells circulating between 1-3.5 km depth

→ Contribution to thermal anomaly in the uppermost part

- Negative thermal anomalies are interpreted as cooling after drilling operations or water injection

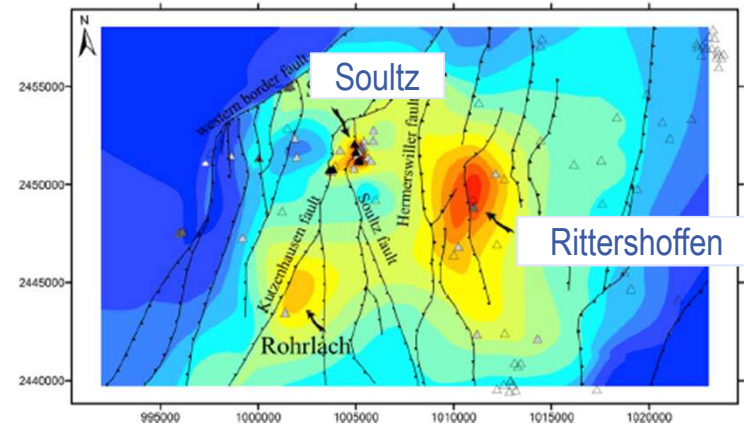
→ Thermal signature of major permeable fracture zones into the wells

Rittershoffen industrial project

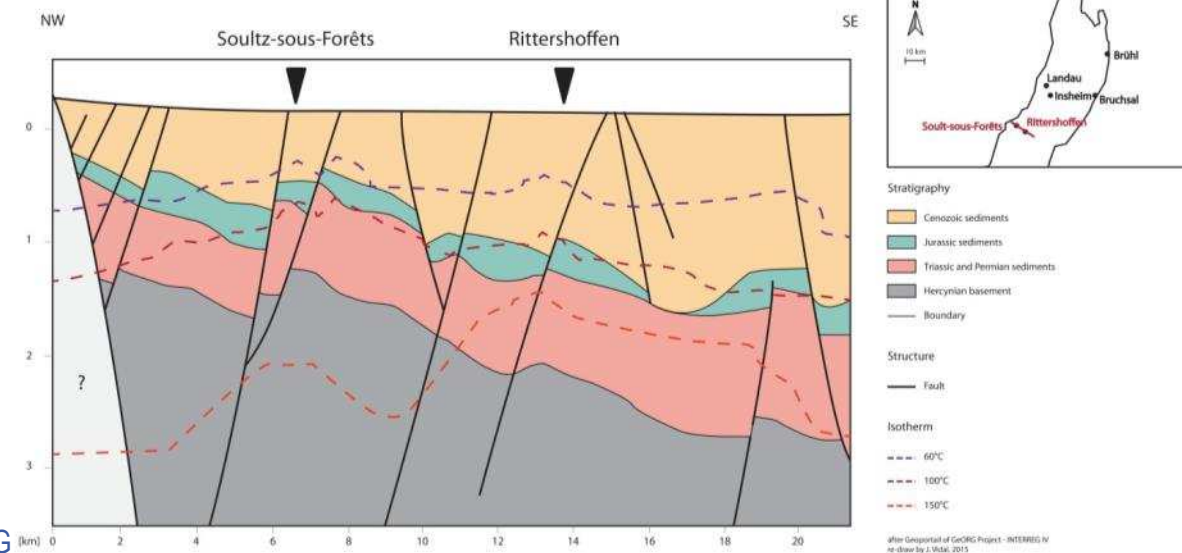
- **2 wells (2.5 - 3 km)**
Expected temperature **170 °C**
Expected flow rate **70 L/s**
Thermal Power **24 MWth**

→ Industrial injectivity reached at shallower depth (<2.5 km depth)

- **Same batholith** than Soutz with an altered and fractured granite at the top of the basement
- **Same thermal profile** than Soutz with convective cells



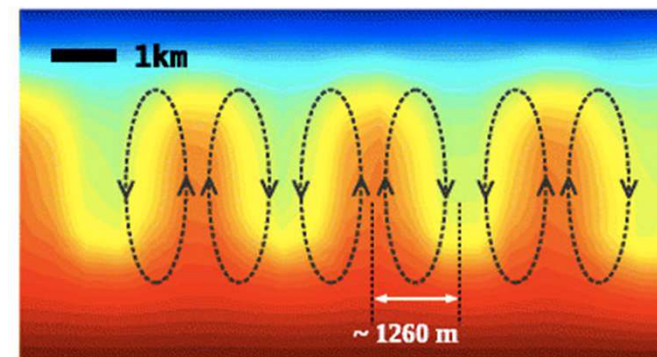
Temperature map (after Bailleux et al., 2013)



Geological cross-section re-draw after GEORG

Conclusion

- Industrial projects are based on lessons learned from Soultz pilot project
- Permeability is higher at the top basement where the granite is hydrothermally altered and fractured
 - **More permeability at shallower depth** = new economic challenge
- Fracture network is multiscale (from microfractures into minerals to large scale fractures)
- Convective cells circulate through fracture network
 - Fracture sealing and deposits decrease permeability



Convective cell modeling (Magnenet et al., 2014)



QUESTIONS?

